

Secondary ice processes: a modulator of global mixed-phase clouds

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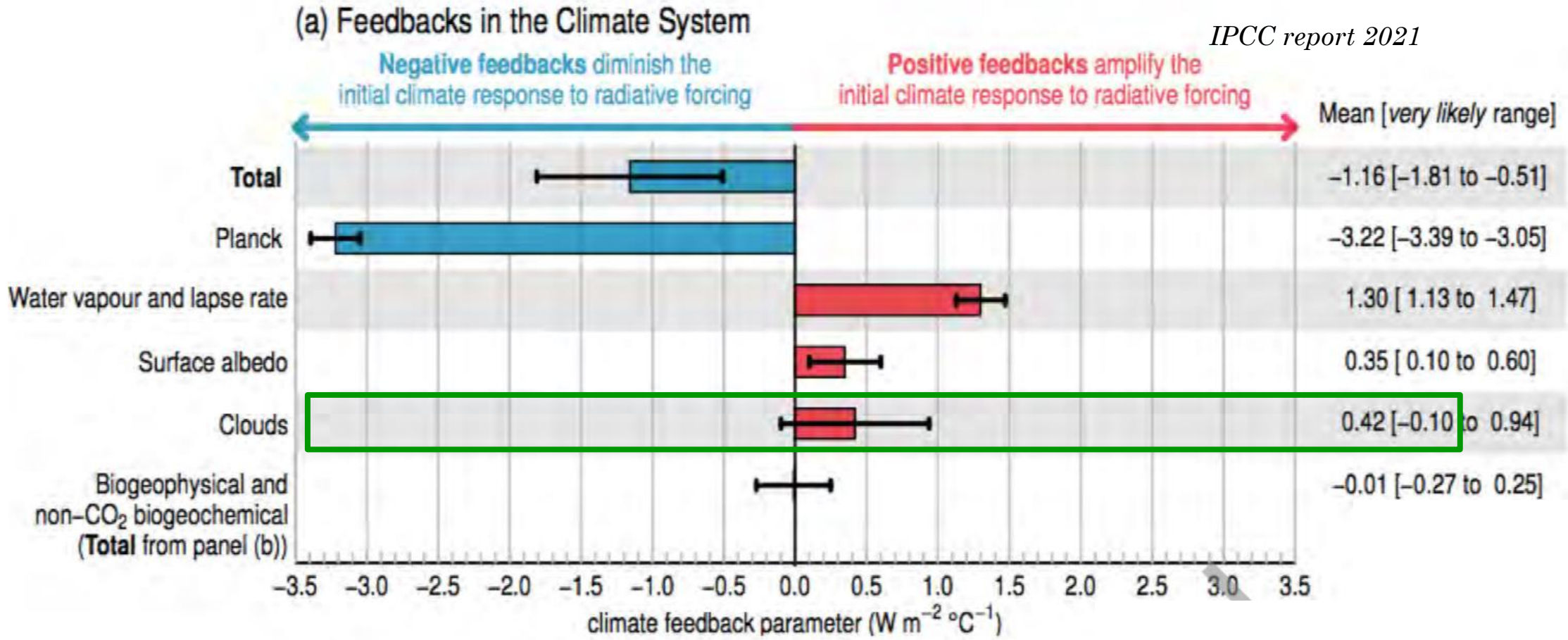
¹Laboratory of Atmospheric Processes and their Impacts, Ecole Polytechnique Federale de Lausanne, Switzerland.

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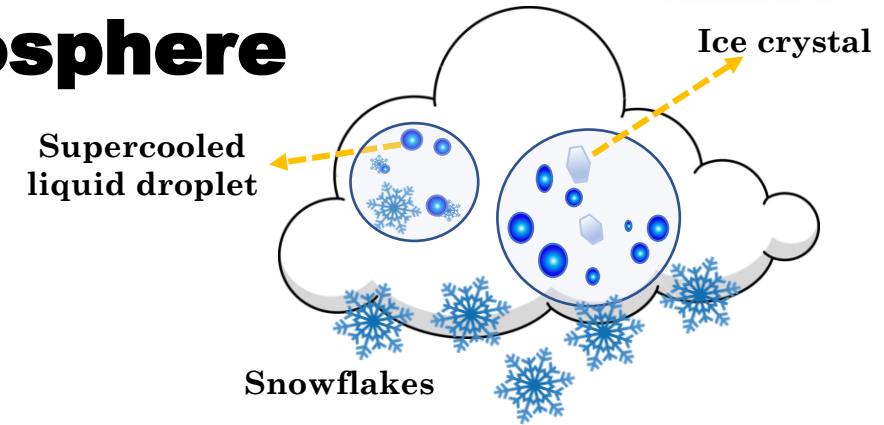
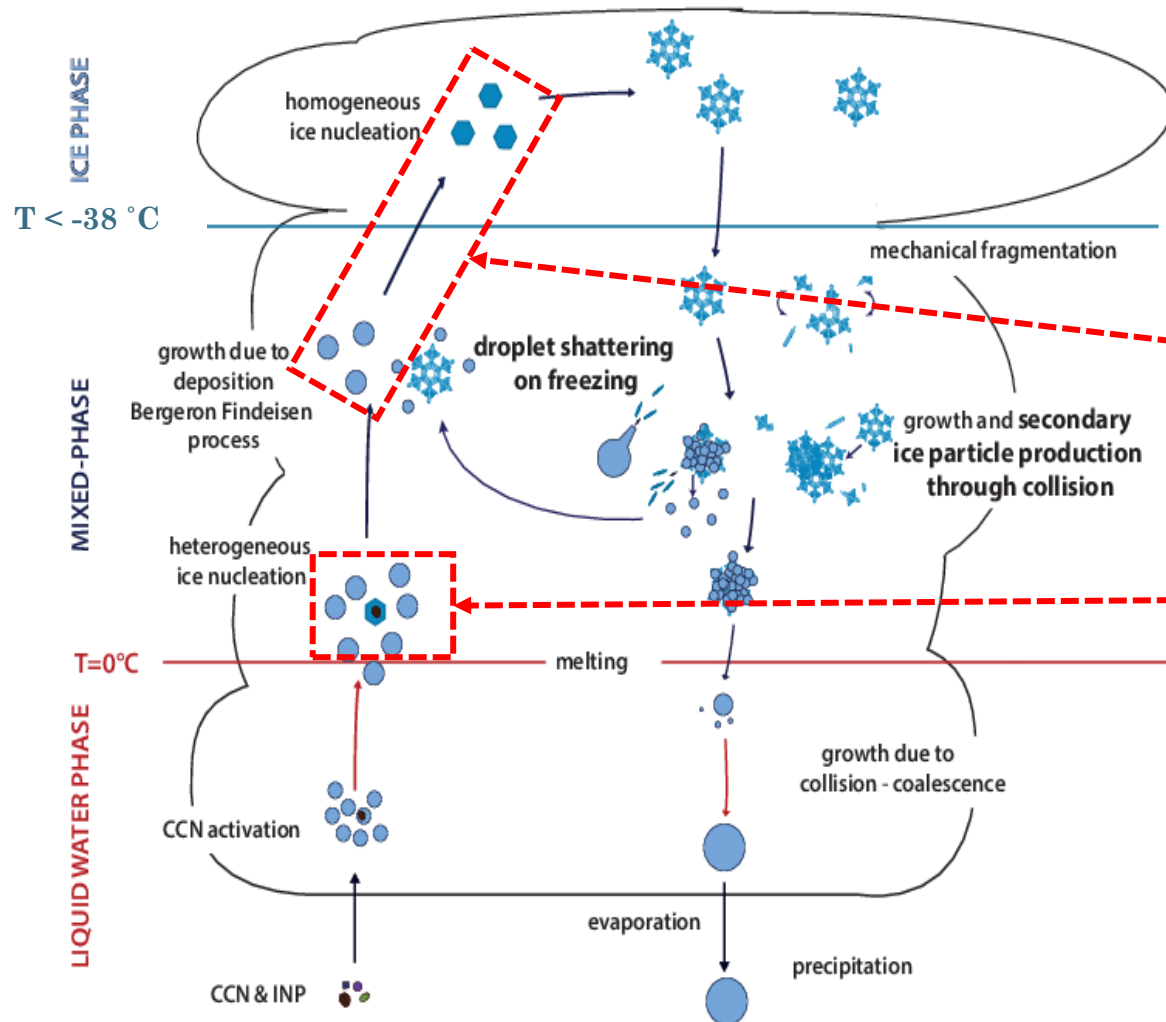
INP Colloquium Talk
December 8, 2022





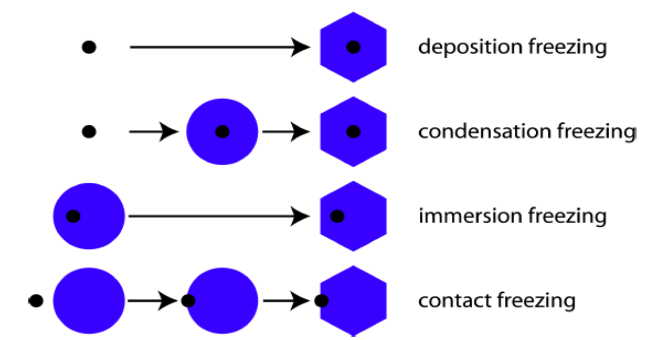
Clouds are the largest source of uncertainty in projections of future climate

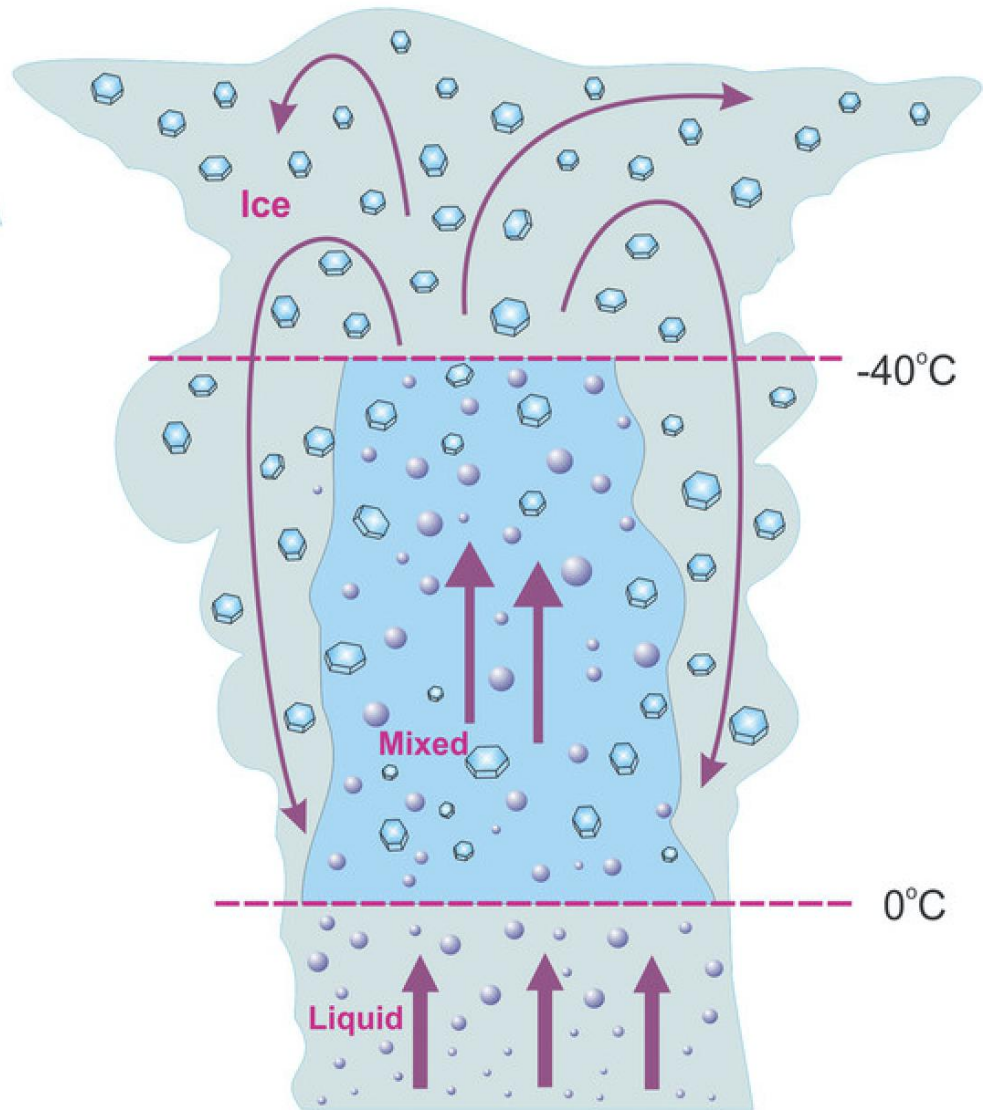
Mixed-phase clouds (MPCs) in the atmosphere



Homogeneous (& Heterogeneous) freezing:
 At $T < -38\text{ °C}$ ice nucleation occurs from the liquid phase (and also with the assistance of INPs).

Heterogeneous freezing:
 Under mixed-phase conditions ($-38\text{ °C} < T < 0\text{ °C}$) the assistance of ice nucleating particles (INPs) is needed to initiate primary ice production

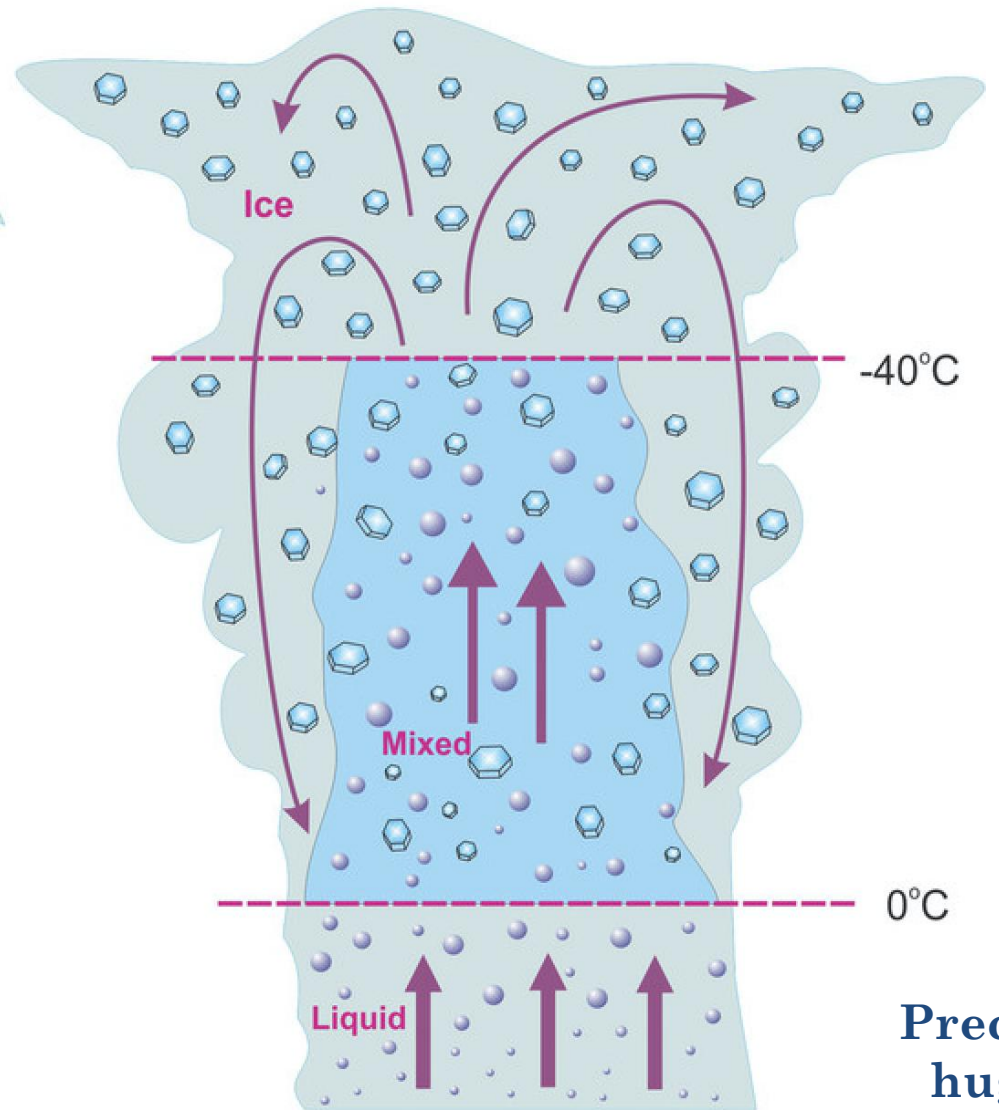




Precipitation at mid- and high-latitudes mostly generated from the mixed- and ice- cloud phase

Mulmenstadt et al . 2015

Mixed-Phase clouds control precipitation



Precipitation at mid- and high-latitudes mostly generated from the mixed- and ice- cloud phase

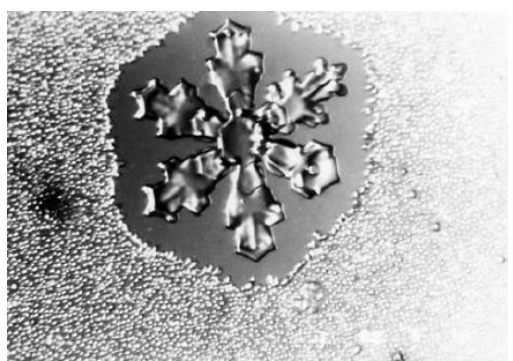
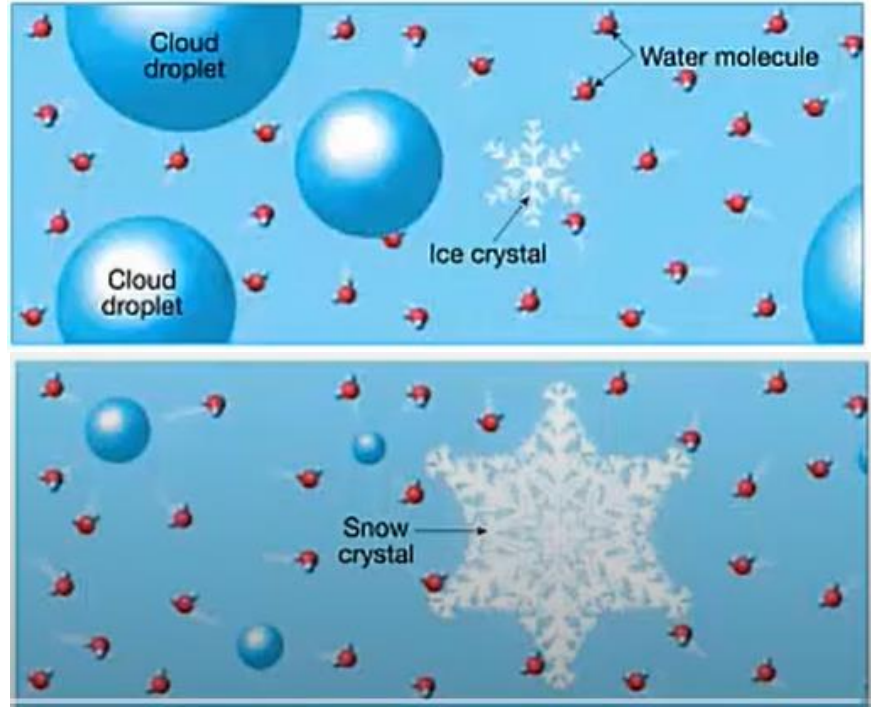
Mulmenstadt et al . 2015

Precipitation extremes have huge impacts on economy and society at large.

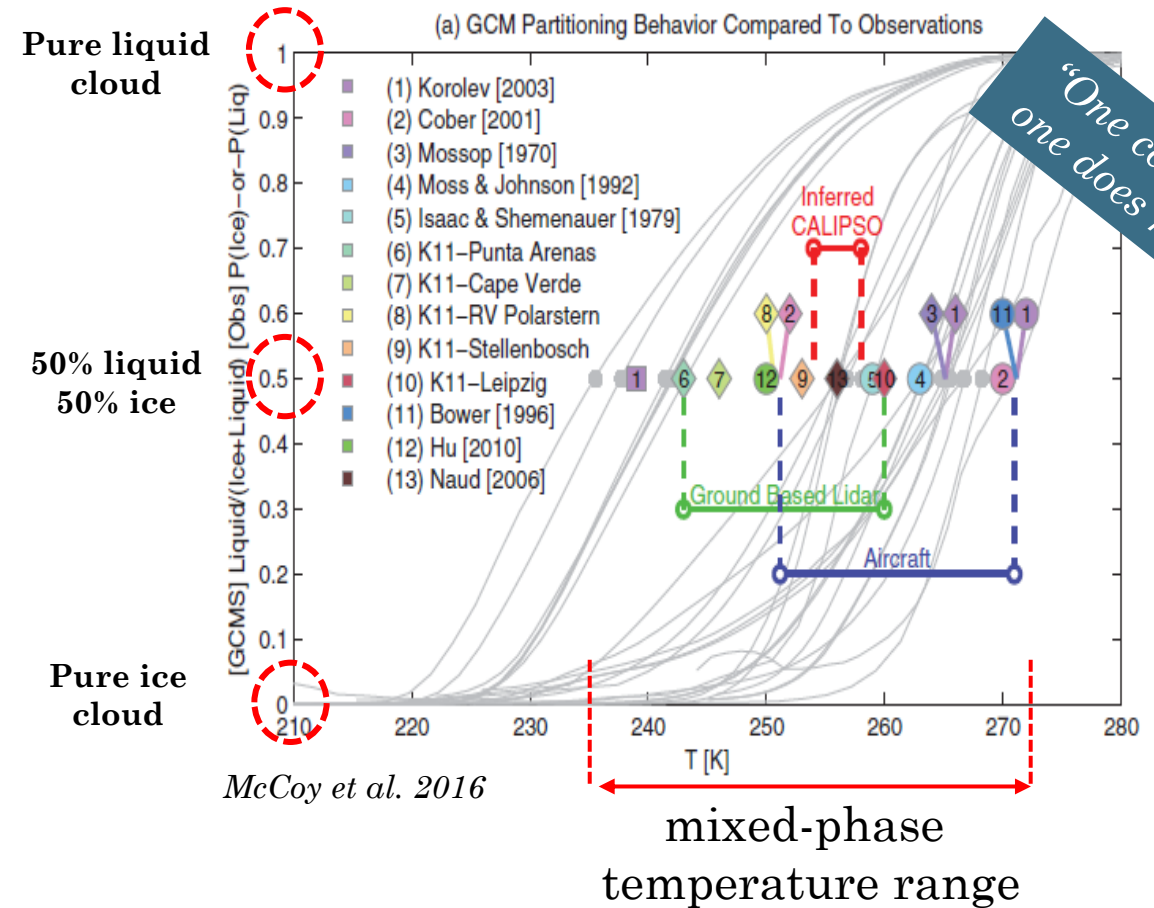


Challenges of representing MPCs within modeling frameworks

- ✓ Important to predict the **amount** and **distribution** of ice and liquid (liquid-ice phase partitioning) in MPCs
- ✓ Models tend to convert water to the ice phase too aggressively



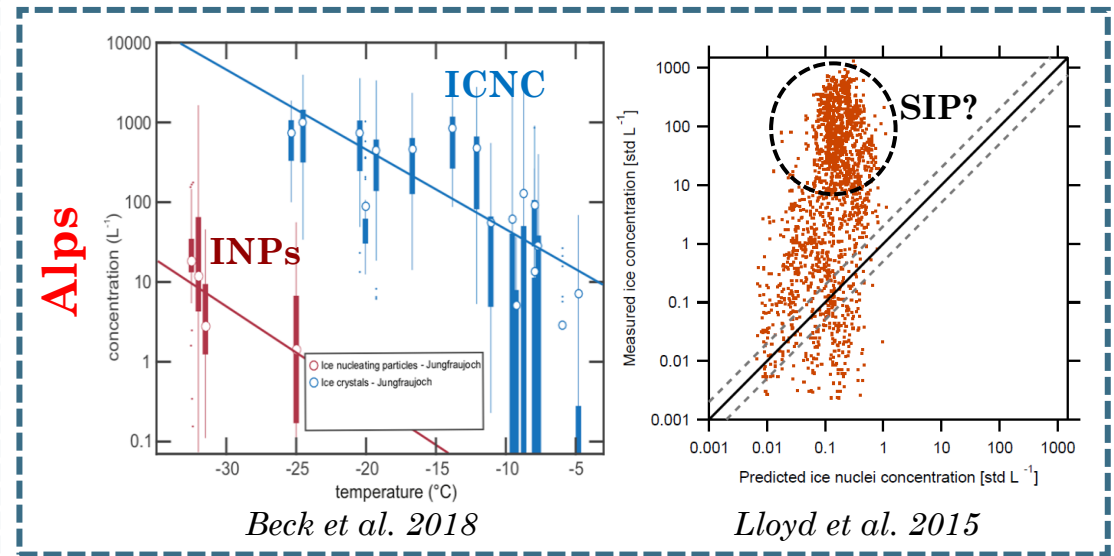
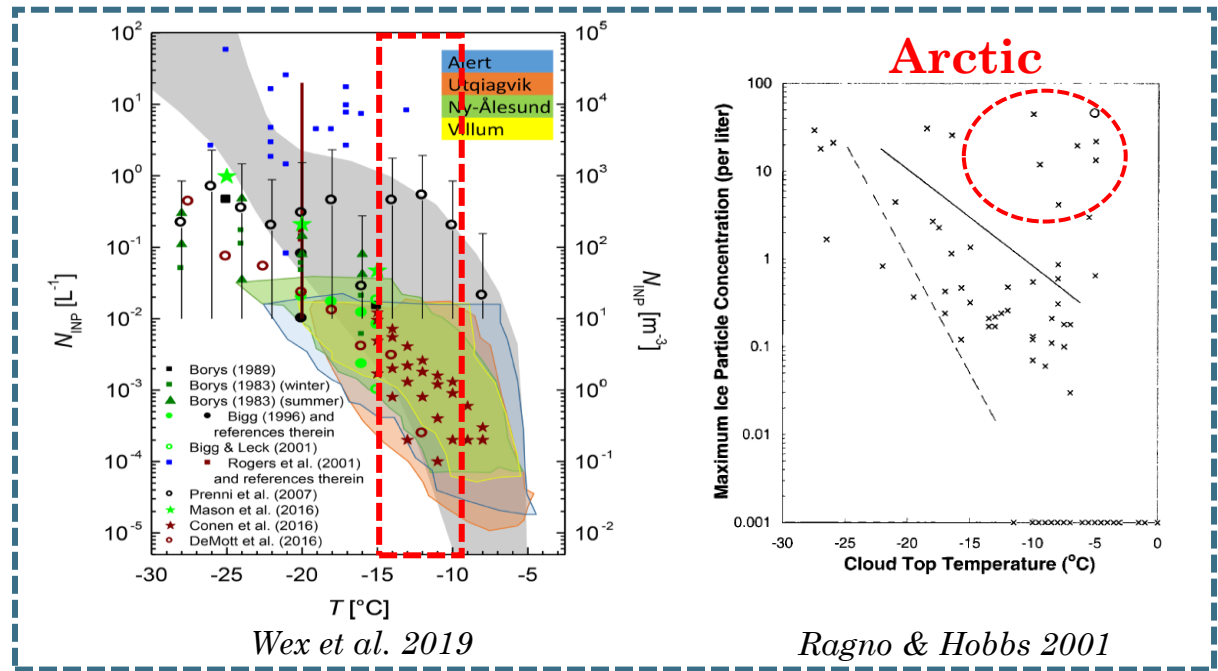
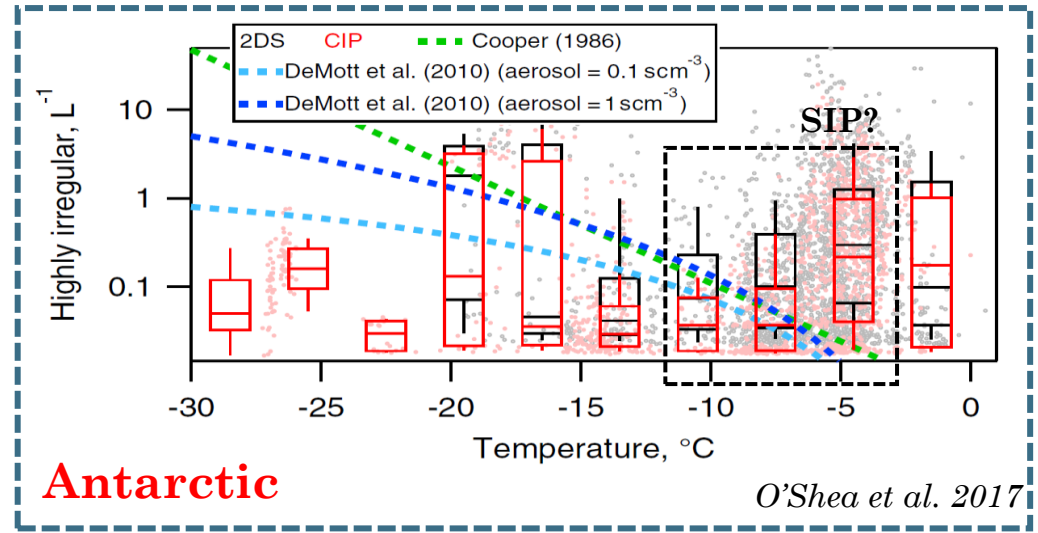
Wegener-Bergeron-Findeisen process (WBF)



McCoy et al. 2016

Measured Ice Crystal Number Concentrations (ICNCs) \gg pre-cloud INPs

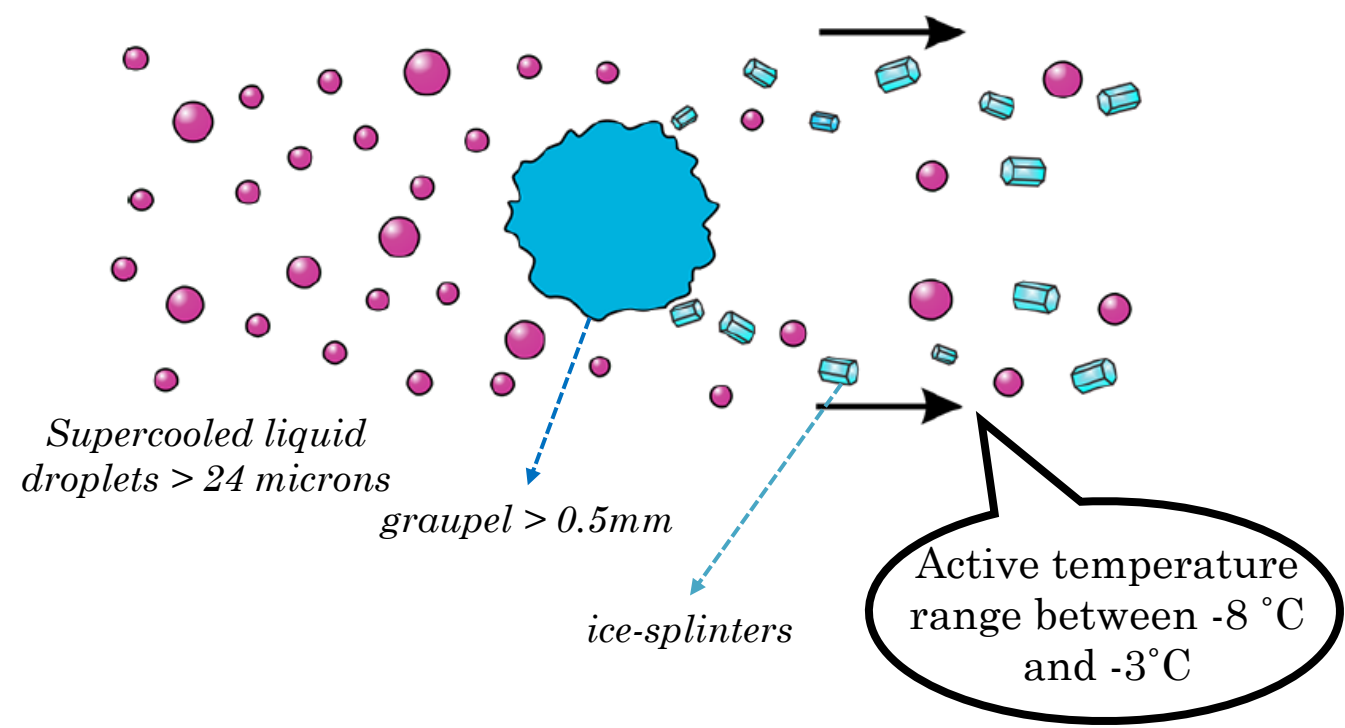
- ✓ Ice Nucleating Particles (INPs) are few in remote polar regions - compared to the ice crystal (ICNCs)
- ✓ Alpine (orographic) clouds have the same behavior.
- ✓ Secondary Ice Production (SIP) processes must be invoked to explain the large difference between INPs and ICNCs



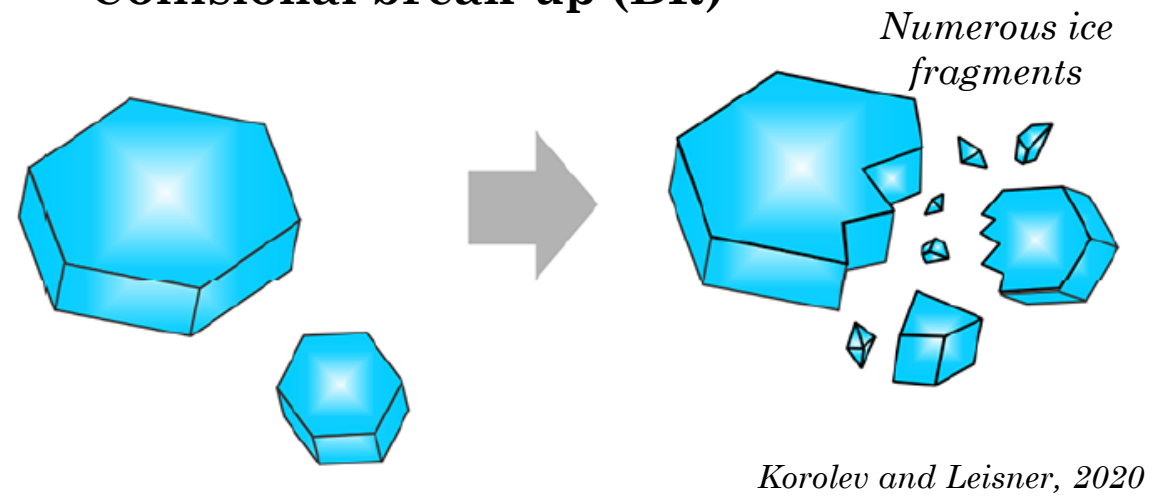
The cause of this cloud-ice paradox → Secondary Ice Production (SIP)*

* SIP = multiplication of primary ice crystals through “other processes” not involving INPs

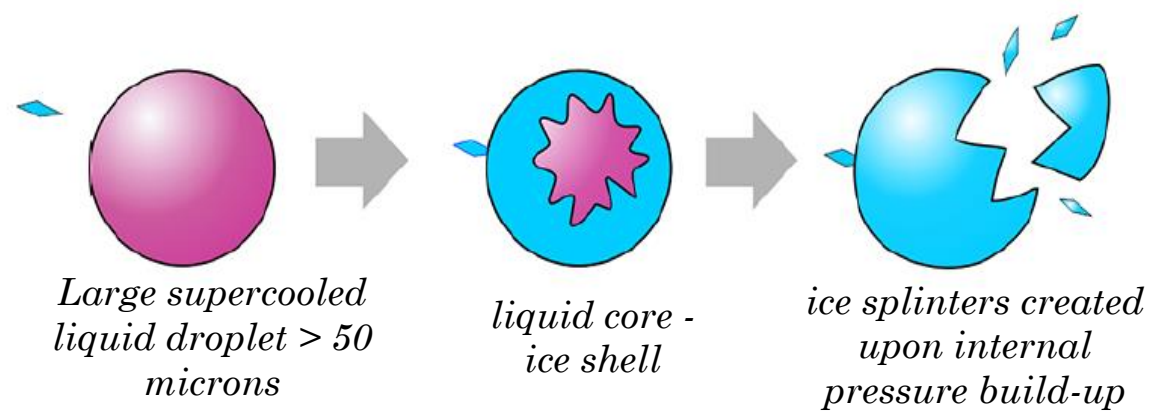
✓ Rime Splintering (RS) or the Hallett-Mossop process (H-M)



✓ Collisional break-up (BR)

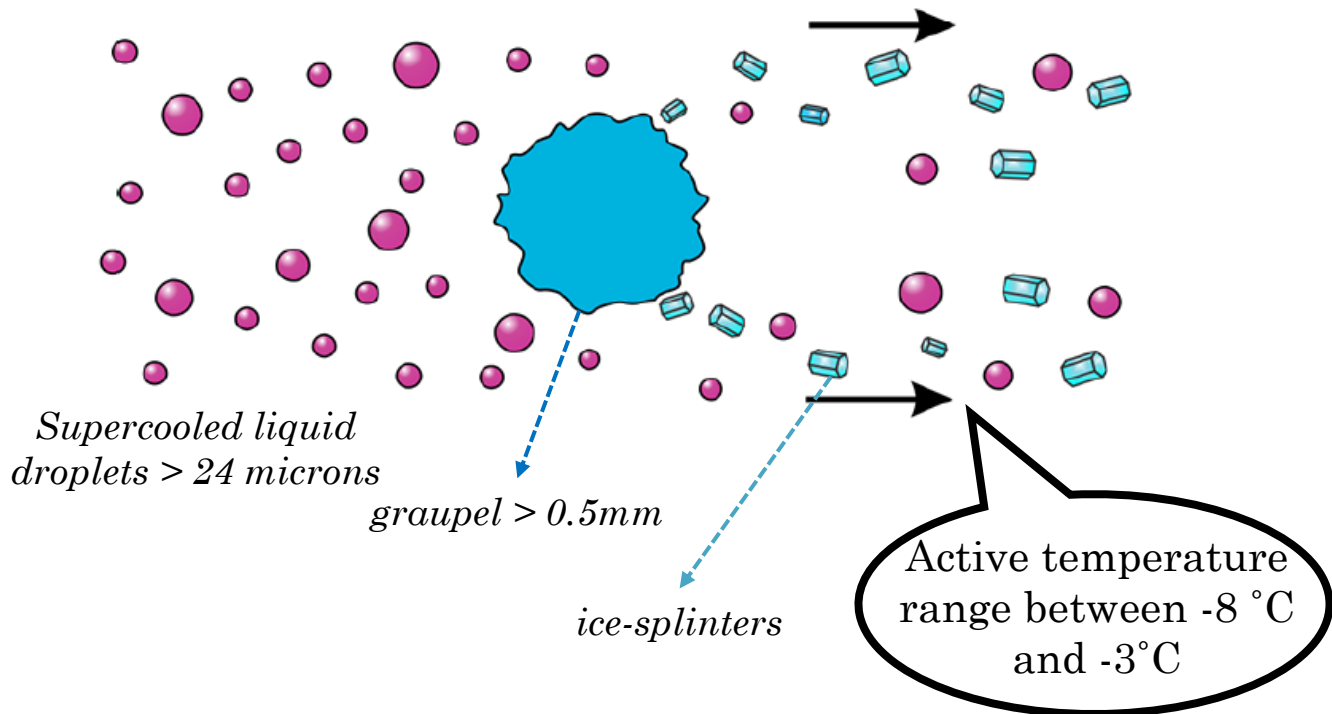


✓ Droplet Shattering (DS) during freezing

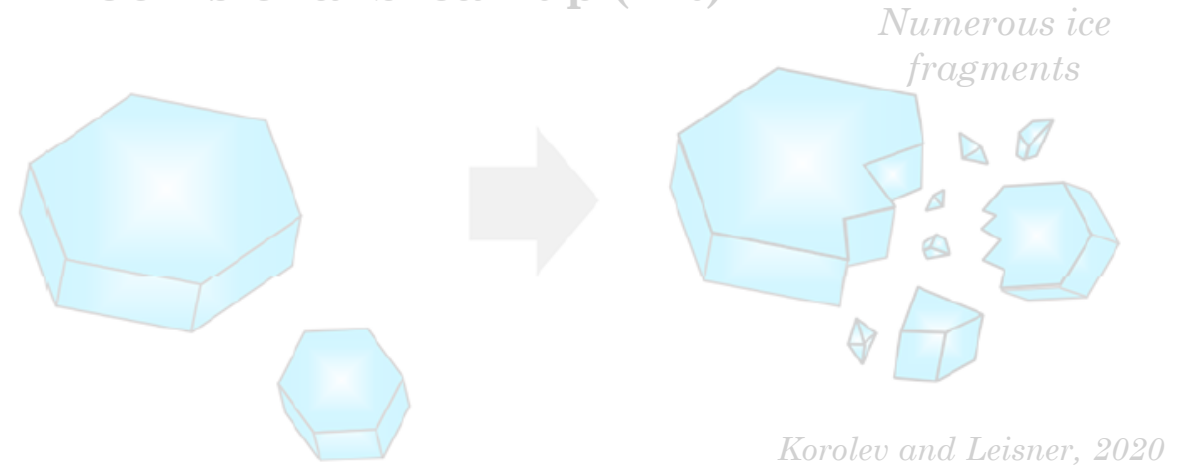


What's included mostly in models:

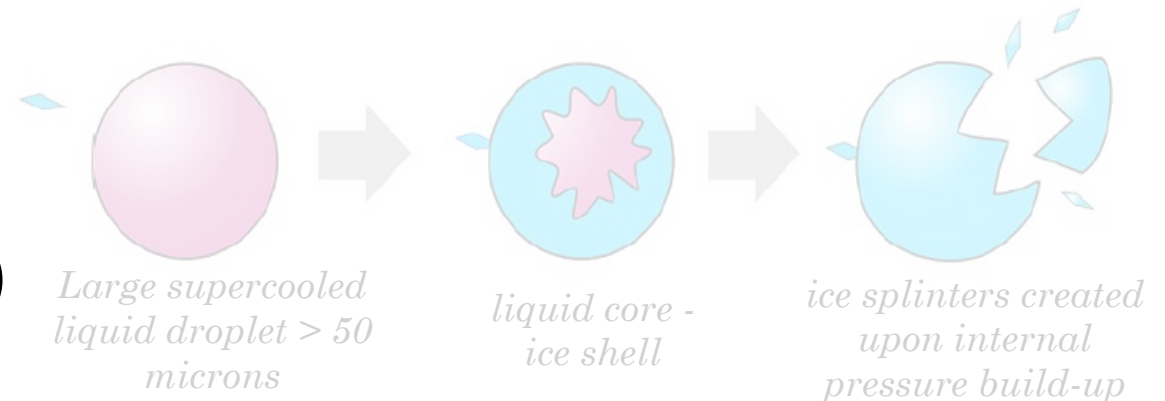
- ✓ **Rime Splintering (RS) or the Hallett-Mossop process (H-M)**



- ✓ **Collisional break-up (BR)**

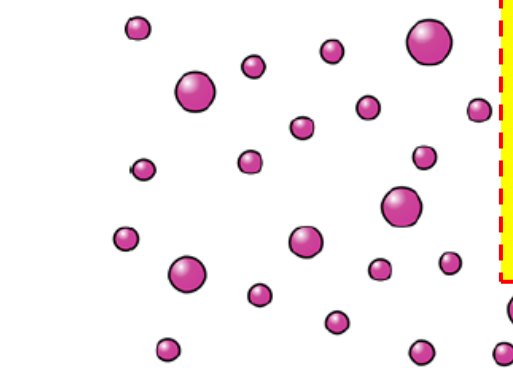


- ✓ **Droplet Shattering (DS) during freezing**



What's included mostly in models:

✓ Rime Splintering process (H-M)



Supercooled liquid droplets > 24 microns

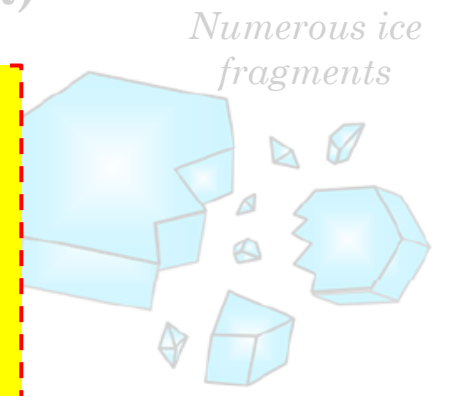
graupel > 0.5mm

ice-splinters

**What happens when we begin including all these processes together?
Let's see for different cloud types and model hierarchies (LES, mesoscale, global).**

Active temperature range between -8 °C and -3°C

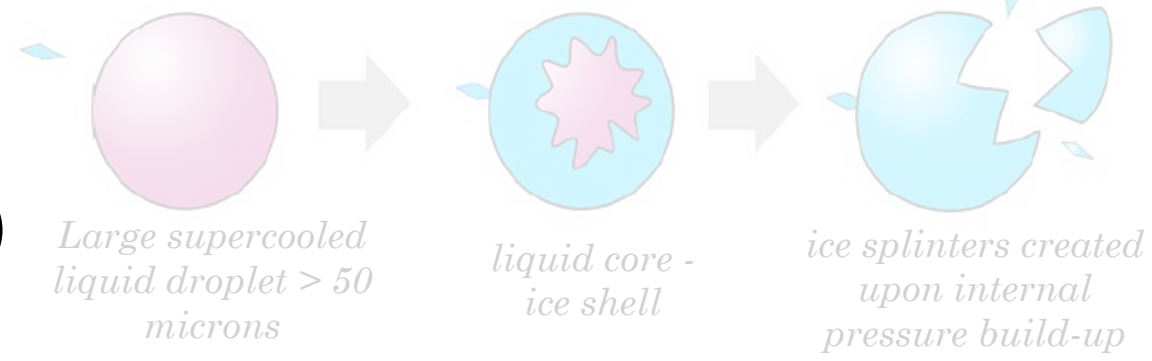
✓ Collisional break-up (BR)



Numerous ice fragments

Korolev and Leisner, 2020

during freezing

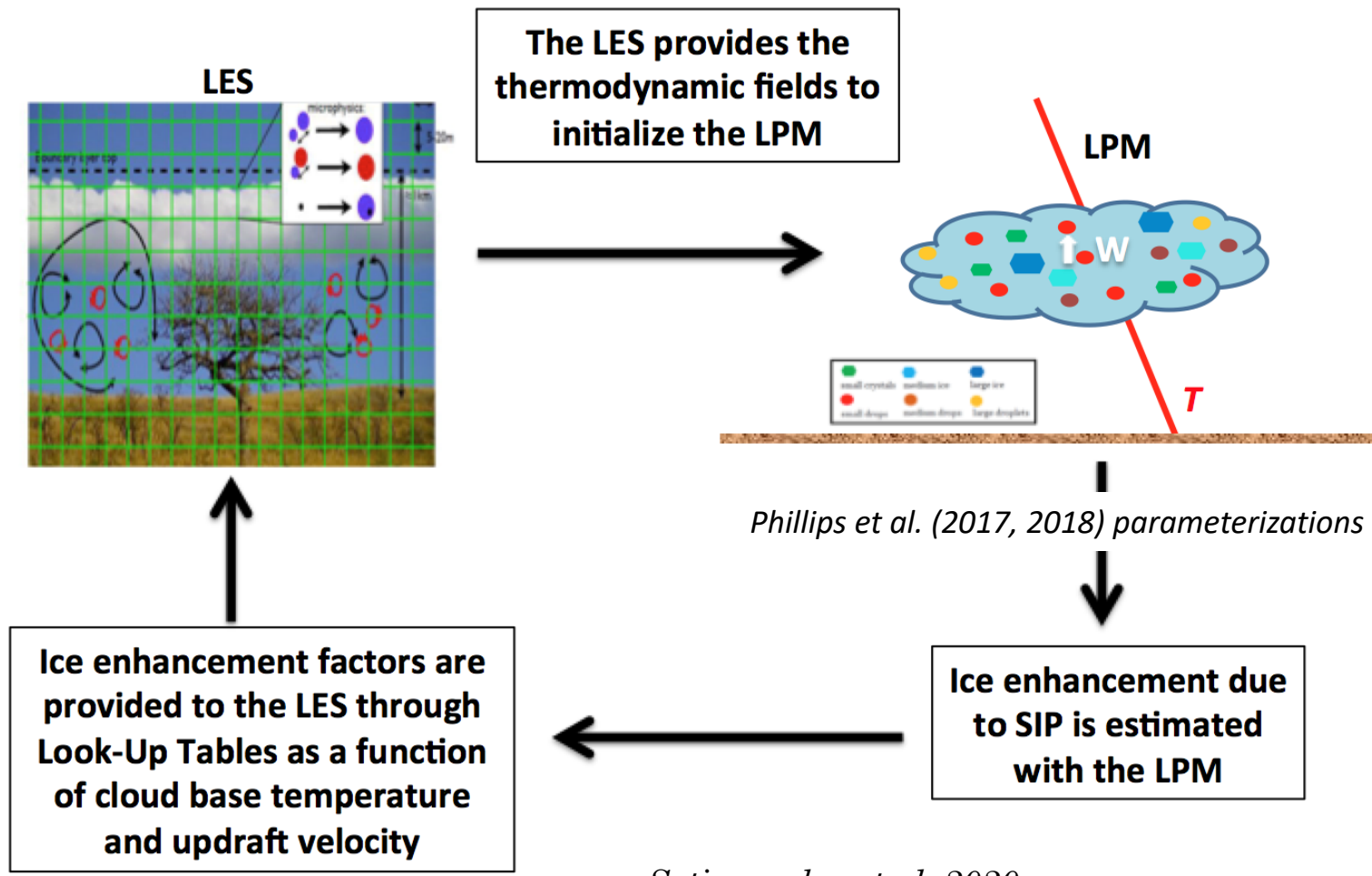


Large supercooled liquid droplet > 50 microns

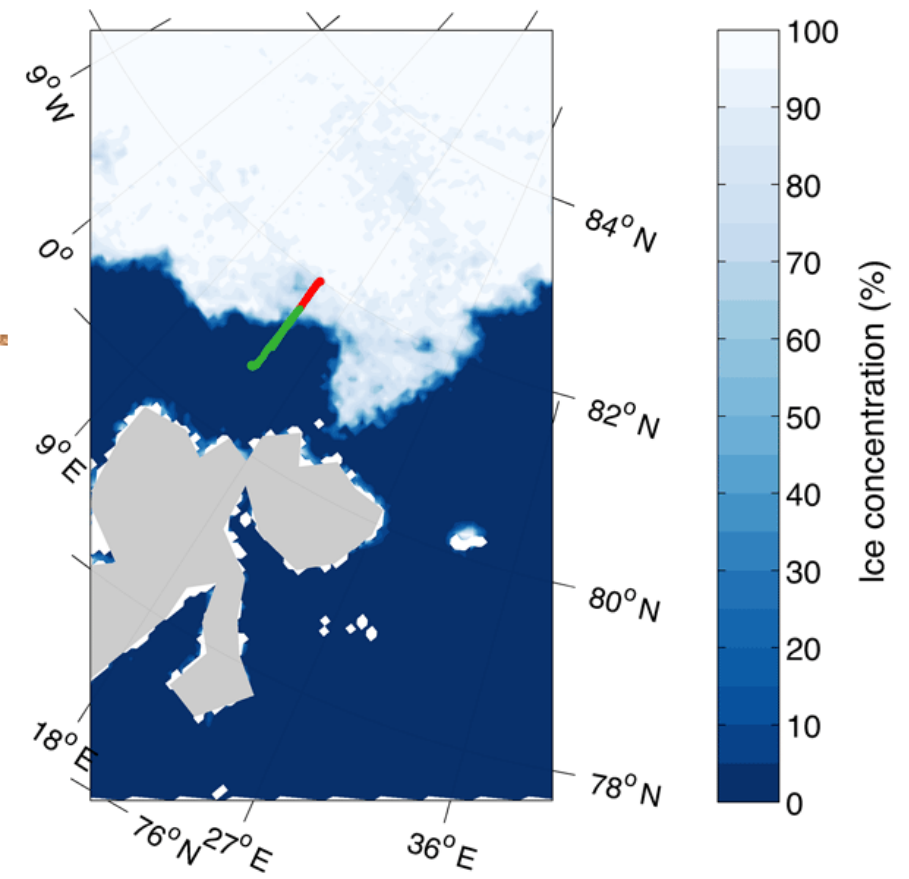
liquid core - ice shell

ice splinters created upon internal pressure build-up

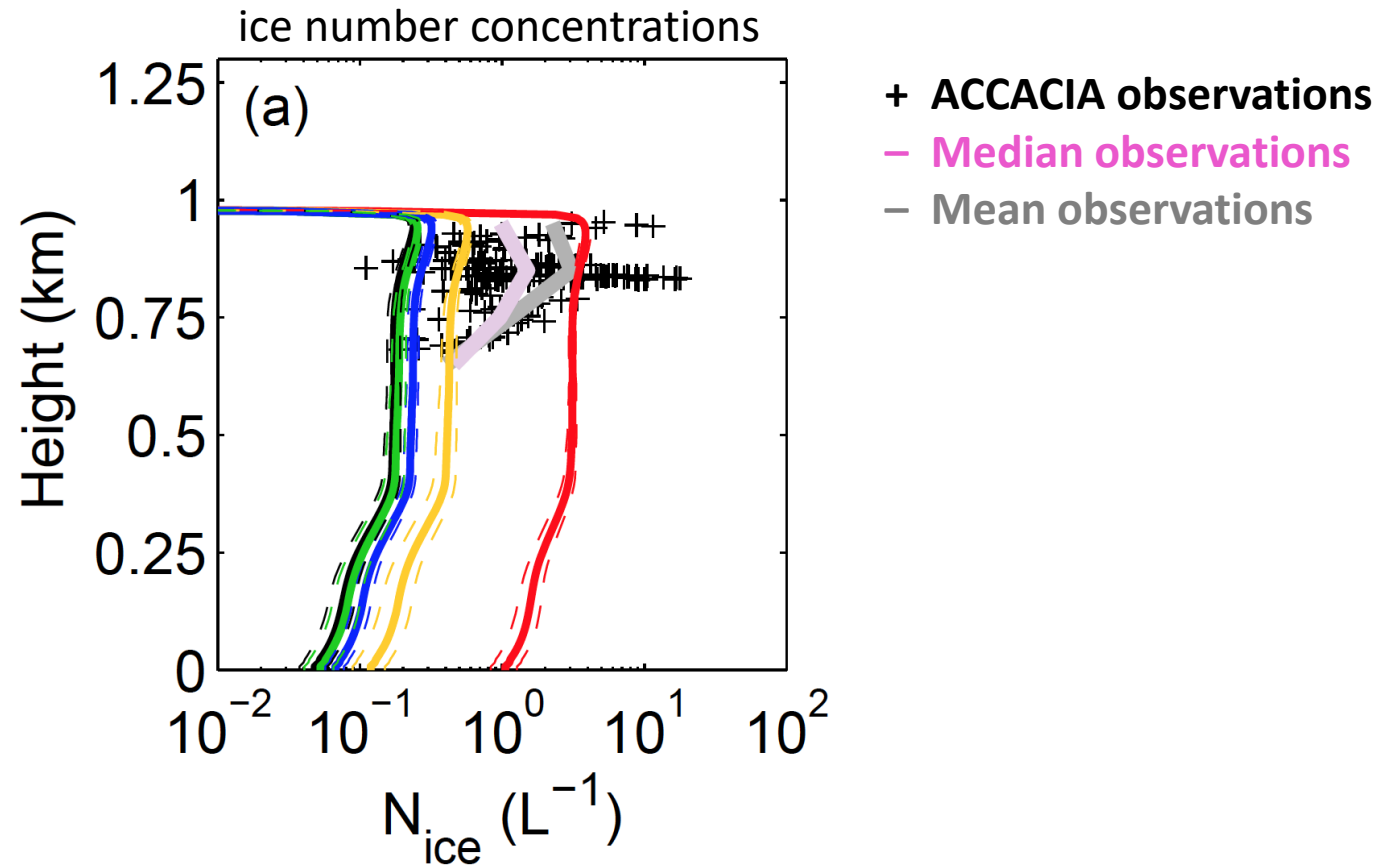
Secondary ice effects in an Arctic cloud deck MIMICA LES with a Lagrangian parcel model (LES-scale)

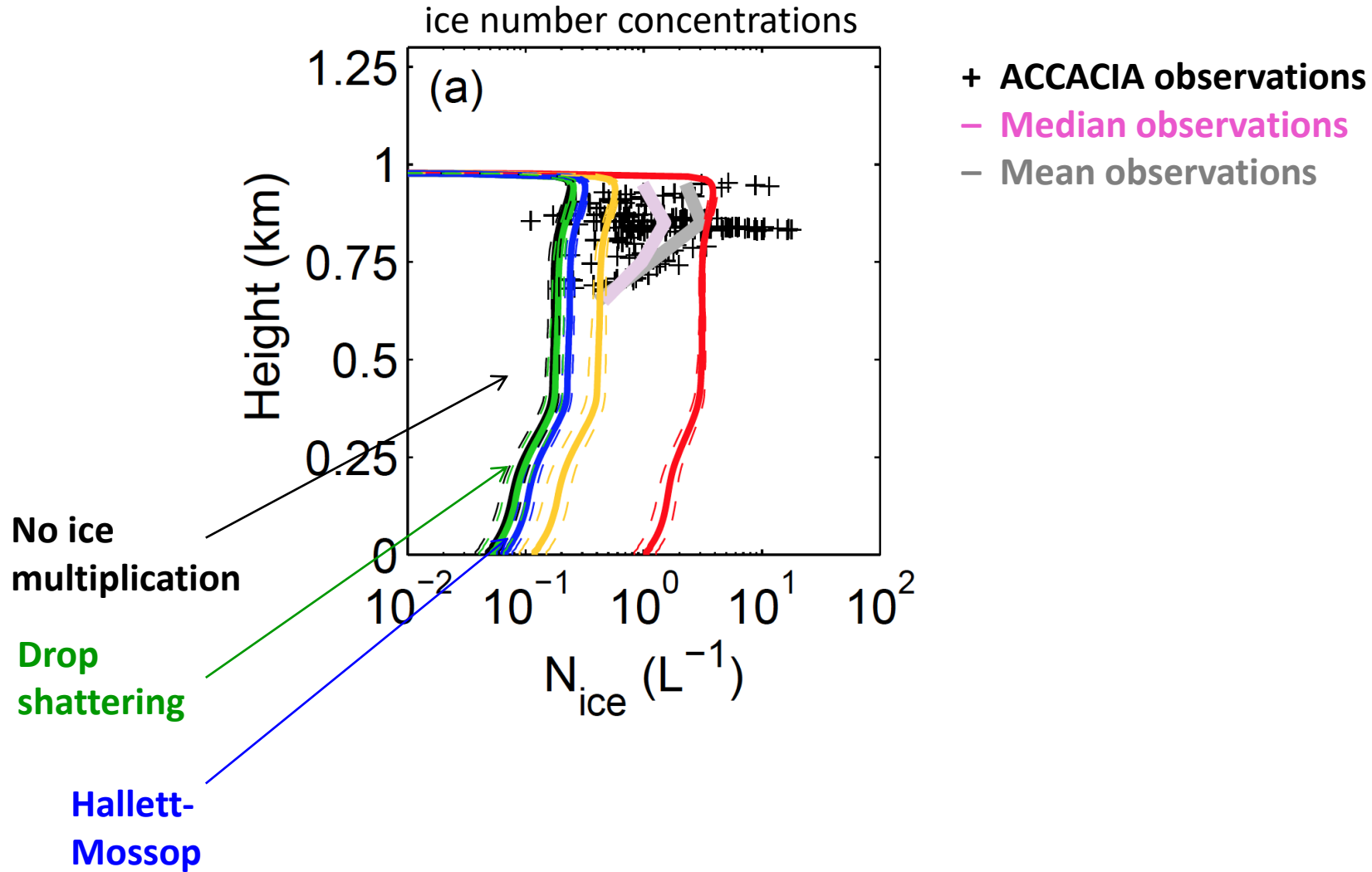


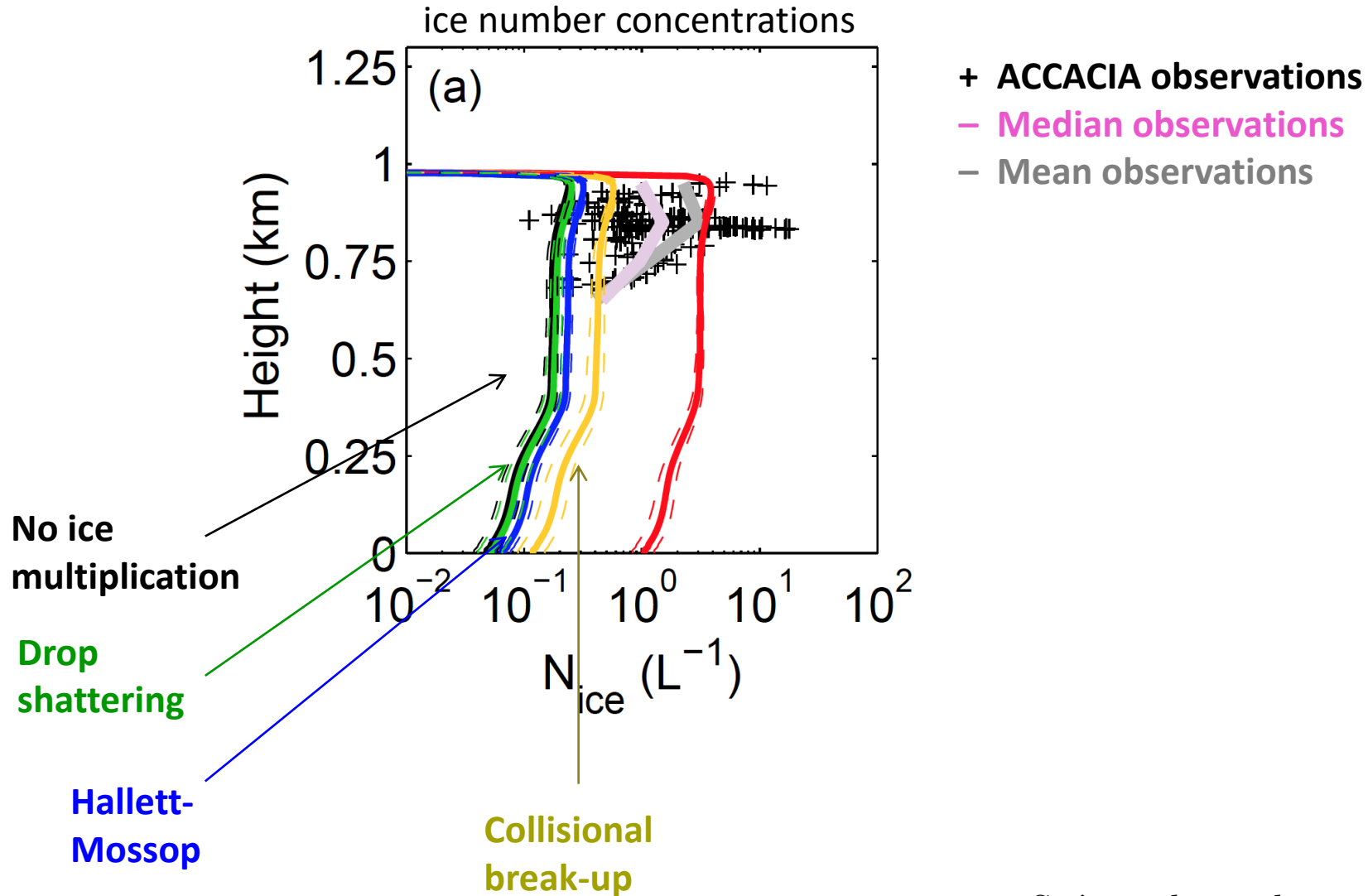
Aerosol-Cloud Coupling and Climate Interactions in the Arctic (ACCACIA) campaign (Svalbard 3,4,7/2013)

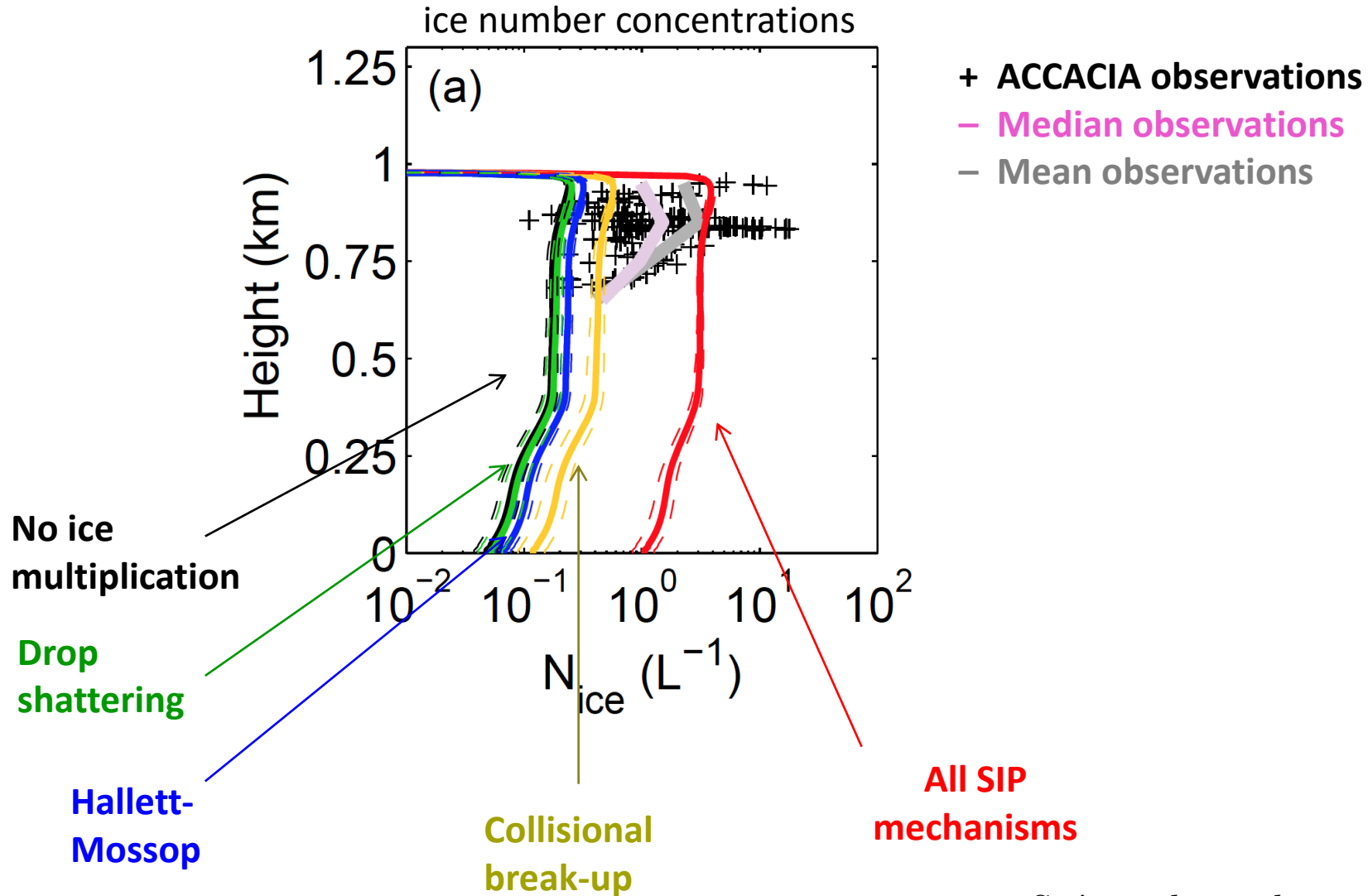


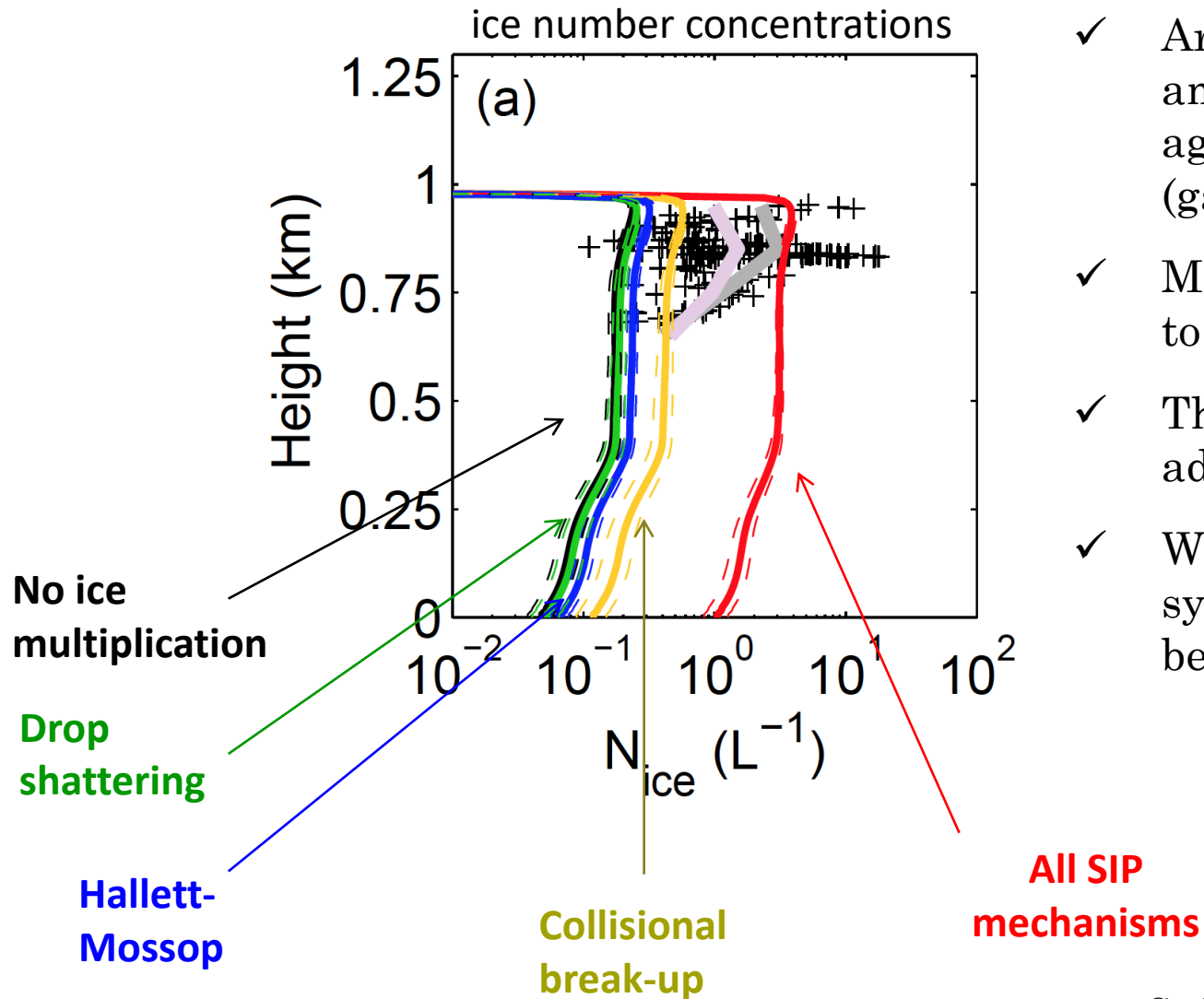
Sotiropoulou et al. 2020







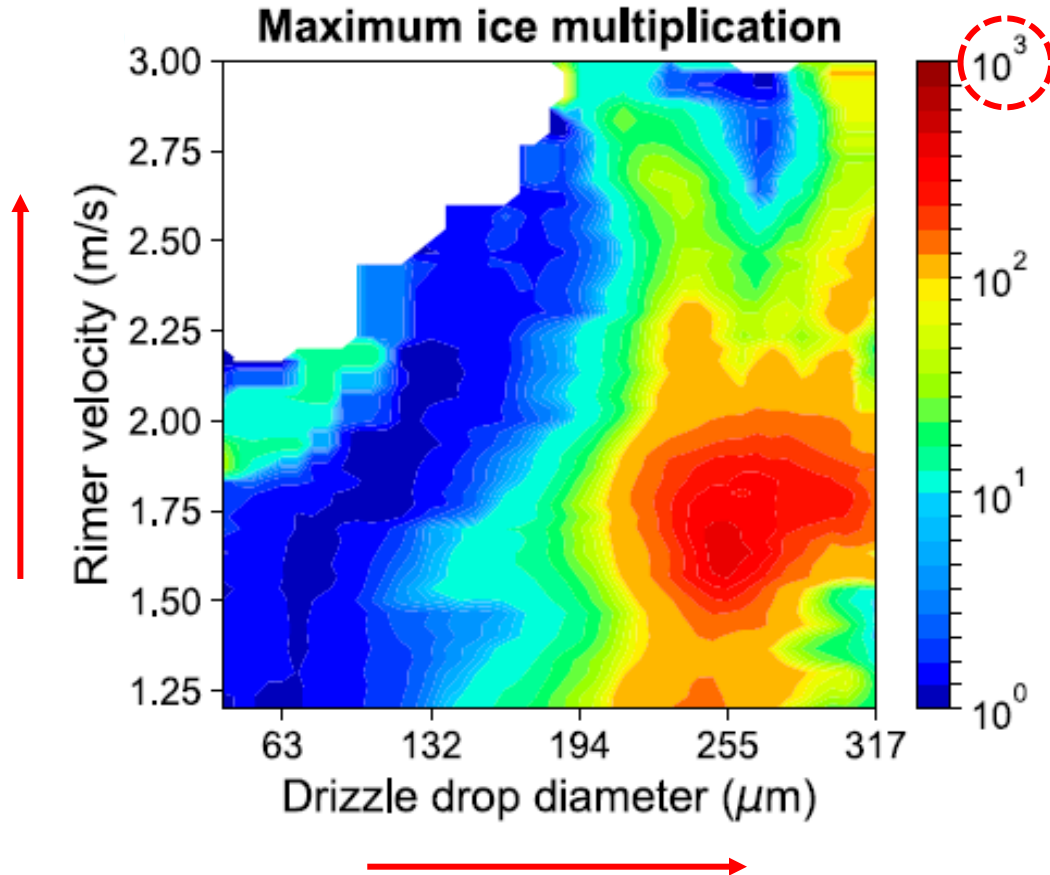




- ✓ Arctic clouds can have considerable amounts of SIP – which goes against established thought (gained from only Hallett-Mossop)
- ✓ Mechanisms can act synergistically to produce ice
- ✓ The effects are not necessarily additive.
- ✓ With many mechanisms active, the system often can exhibit buffered behavior (not shown).

Droplet shattering can be important as well

- ✓ 6 years of **cloud radar** data of slightly supercooled Arctic clouds ($-10\text{ °C} < T < 0\text{ °C}$) in Utqiagvik (Barrow), Alaska (2013 – 2019)
- ✓ Relative frequency of SIP <10%
- ✓ Even if rare, our models don't reproduce this behavior means there is improvement needed in getting the big droplets there.



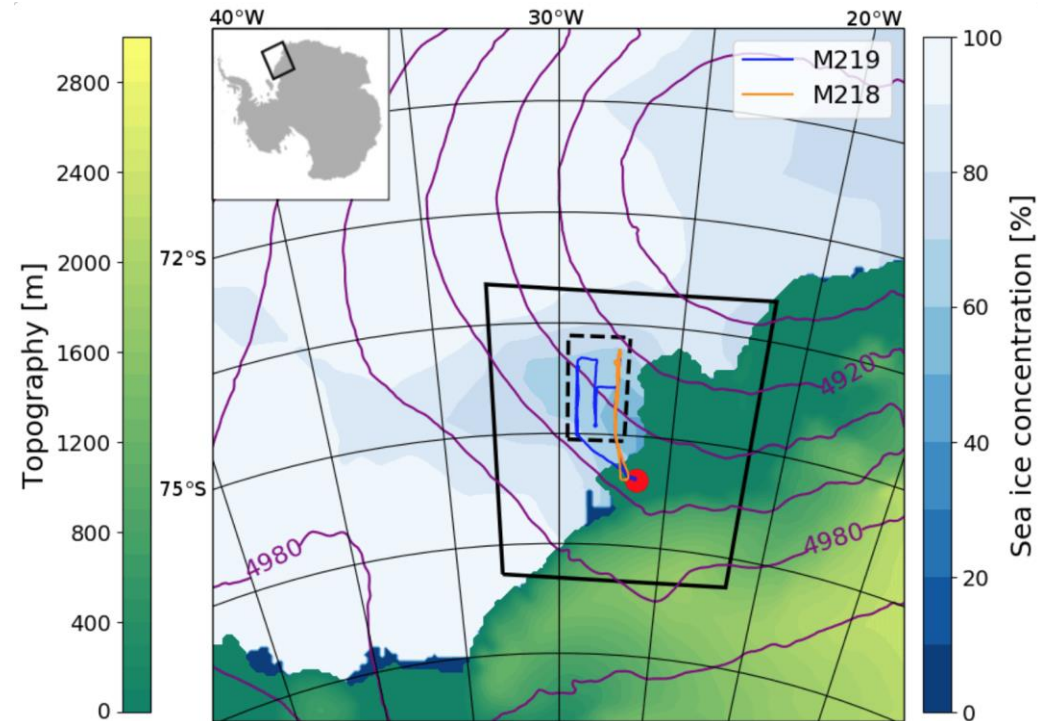
Quantification of the **ice-enhancement factor** (i.e., the enhancement in ice number concentration due to SIP)

Luke et al., 2021, PNAS

Conditions with large drizzle drops can lead to explosive SIP!

Secondary ice effects in Antarctic clouds

WRF (Weather Research & Forecasting model) (regional scale)



Simulations of 2 MAC (Microphysics of Antarctic clouds campaign) flights.
November–December 2015 over coastal Antarctica and the Weddell Sea

Collisional Break-up

Morrison scheme: 2-moment bulk microphysics, 5 hydrometeor species:



Fragmentation is assumed to occur after:

- (1) cloud ice – graupel collisions → fragmentation of ice
- (2) cloud ice – snow collisions → fragmentation of ice
- (3) snow – graupel collisions → fragmentation of snow
- (4) snow – snow collisions → fragmentation of snow
- (5) graupel – graupel collisions → fragmentation of graupel

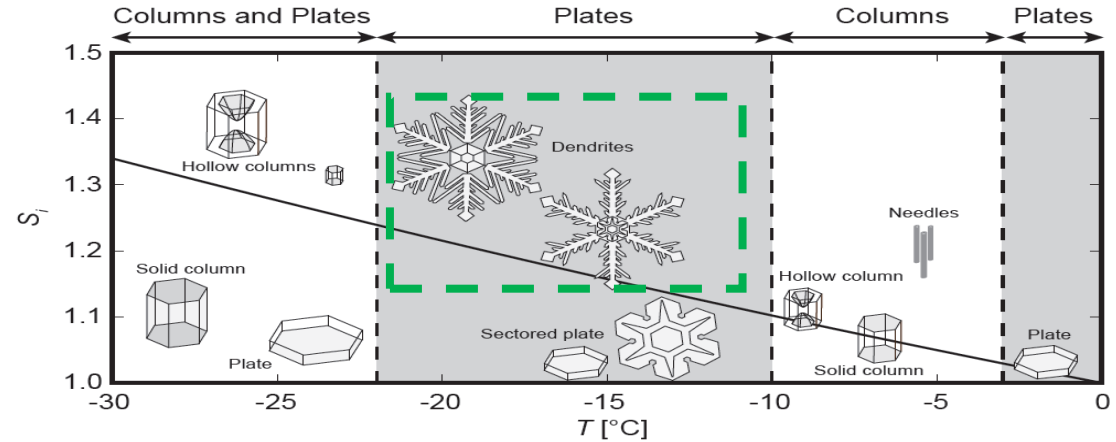


Fragments are added to **cloud ice** category

Fragments per collision: parameterizations

I. Following Phillips et al. (2017)

$$F_{BR} = aA \left(1 - \exp \left\{ - \left[\frac{C K_o}{aA} \right]^\gamma \right\} \right), \quad a = \pi D^2$$



F_{BR} is a function of (i) collisional kinetic energy, (ii) size, (iii) ice type, (iv) ice habit and (v) rimed fraction

Parameter	Collision Type			
	Type I: Collisions of graupel (size: $5 \times 10^{-4} < D < 5 \times 10^{-3}$ m) with graupel/hail	Type I: Collisions between hail only	Types II or III: Collisions of either crystals or snow (rimed fraction: $\Psi < 0.5$; size: $5 \times 10^{-4} < D < 5 \times 10^{-3}$ m) with any ice (crystals, snow, or graupel/hail)	
			Dendrites (e.g., -12° to -17°C)	Spatial planar (e.g., -40° to -17°C , -9° to -12°C)
A (m^{-2})	$\frac{a_0}{3} + \max\left(\frac{2a_0}{3} - \frac{a_0}{9} T - T_0 , 0\right)$	$\frac{a_0}{3} + \max\left(\frac{2a_0}{3} - \frac{a_0}{9} T - T_0 , 0\right)$	$(1.41 \times 10^6) \times (1 + 100\Psi^2) \times \left(1 + \frac{3.98 \times 10^{-5}}{D^{1.5}}\right)$	$1.58 \times 10^7 \times (1 + 100\Psi^2) \times \left(1 + \frac{1.33 \times 10^{-4}}{D^{1.5}}\right)$
C (J^{-1})	$(6.30 \times 10^6) \times \psi$	3.31×10^5	$(3.09 \times 10^6) \times \psi$	$(7.08 \times 10^6) \times \psi$
γ	0.3	0.54	$0.50 - 0.25\Psi$	$0.50 - 0.25\Psi$
a_0 (m^{-2})	$(3.78 \times 10^4) \times \left(1 + \frac{0.0079}{D^{1.5}}\right)$	4.35×10^5	—	—
T_0 ($^\circ\text{C}$)	-15	-15	—	—
N_{\max}	100	1000	100	100
ψ	3.5×10^{-3}	3.5×10^{-3}	3.5×10^{-3}	3.5×10^{-3}
ζ	0.001	10^{-6}	0.001	0.001



Fragments per collision: parameterizations

II. Following Sullivan et al. (2018)

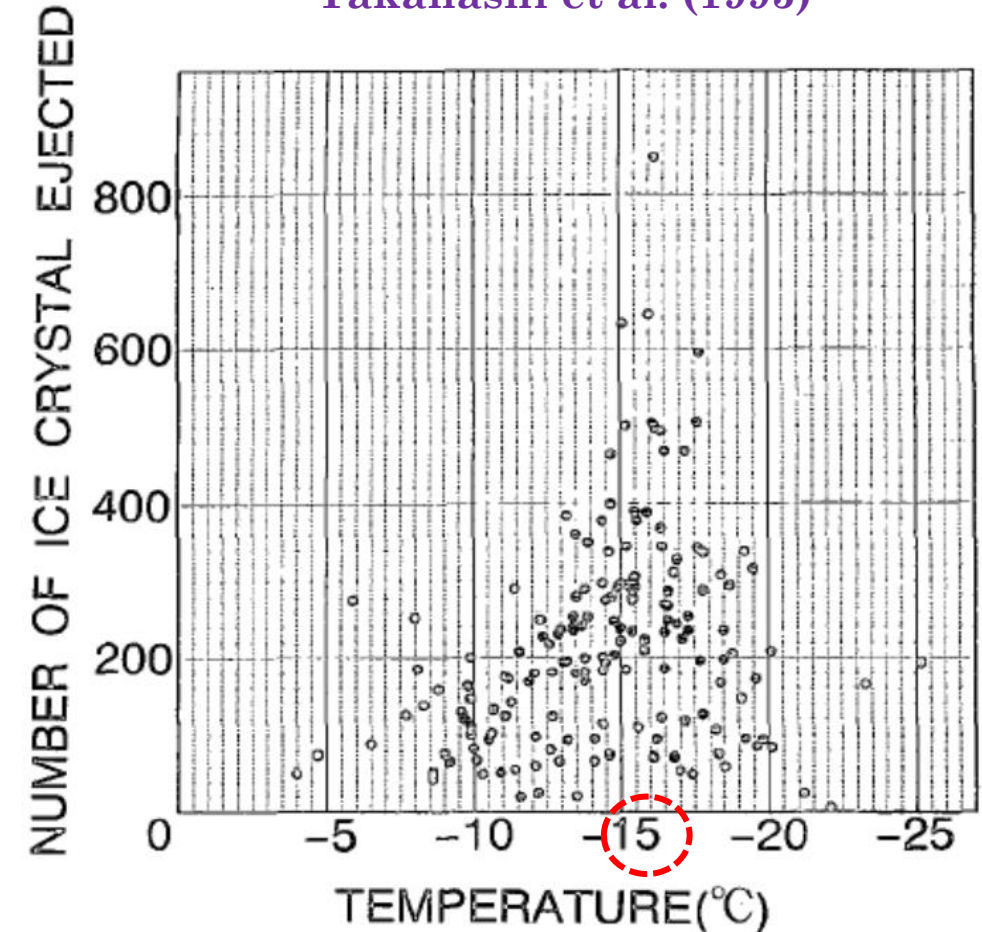
$$F_{BR} = 280(T - 252)^{1.2} \exp\left(-\frac{(T-252)}{5}\right) \frac{D}{D_0}$$

F_{BR} is a function of (i) **temperature** and is further scaled to include the influence of (ii) **size**

Note! Takahashi et al. (1995) used 2-cm hailballs in their experiments ($D_0 = 0.02$ m)

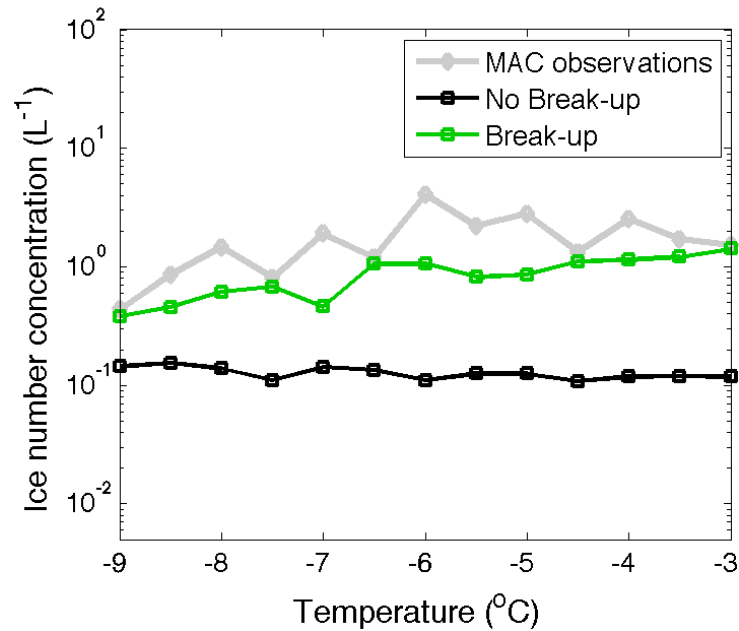
max ice splinter production rate at ~ -16 °C

Laboratory experiment by
Takahashi et al. (1995)

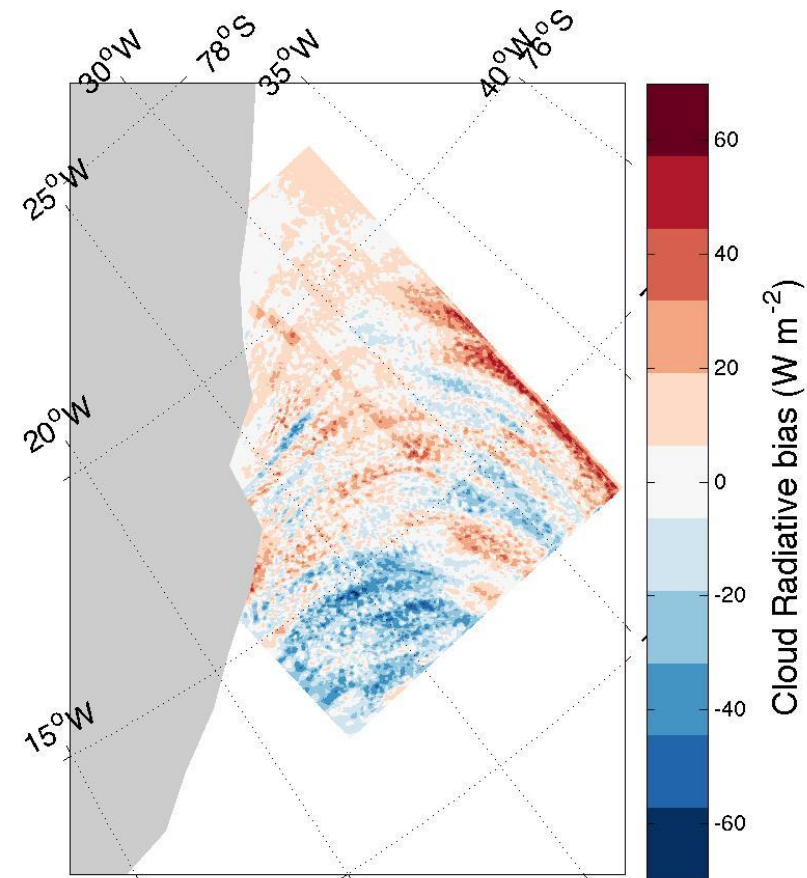


SIP impacts on Antarctic clouds (regional scale)

MAC (Microphysics of Antarctic clouds campaign)



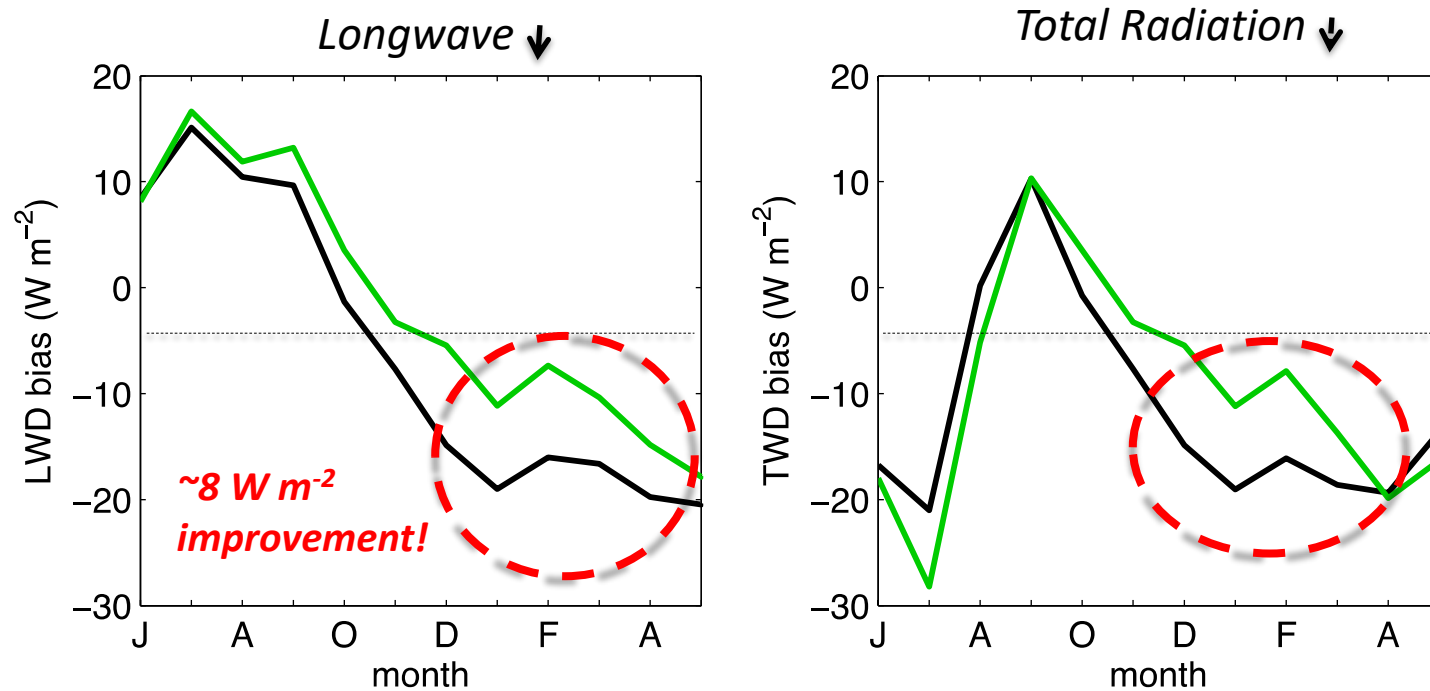
Implementation of the missing mechanism bridges the gap between observed and modeled ice number concentrations



The newly implemented process alters cloud-induced warming by up to 60 W m⁻²

SIP effects in the NorESM2 climate model (global scale)

Radiation biases (model –EBAF satellite observations) over the Arctic region (2016-2017)



- standard* model (only 1 SIP mechanism: Hallett-Mossop)
- additional SIP mechanisms (BR, DS)

Sotiropoulou et al., J. Clim., in review

*ice aggregation adjusted for Arctic clouds following Chellini et al. (2022)

CLACE 2014 field campaign

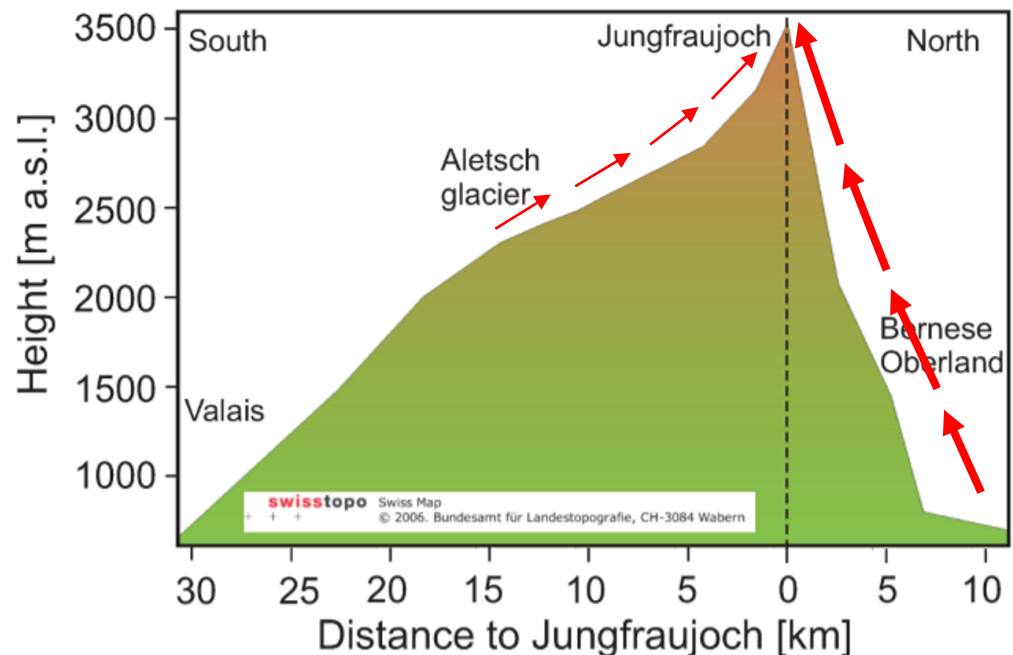
“Cloud Aerosol Characterization Experiments” (Lloyd et al. 2015)



Lohmann et al., 2016

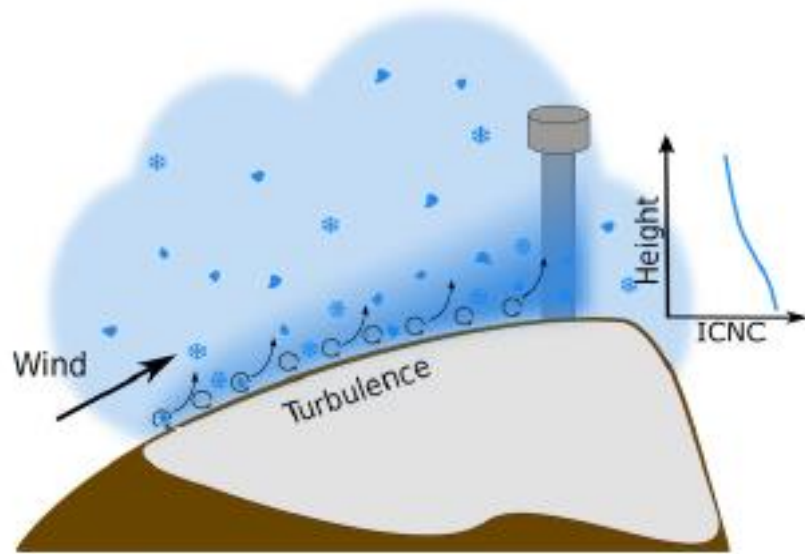
SE winds

- ✓ Gentle ascent
- ✓ Weak updrafts
- ✓ Low supersaturations
- ✓ Glaciated clouds

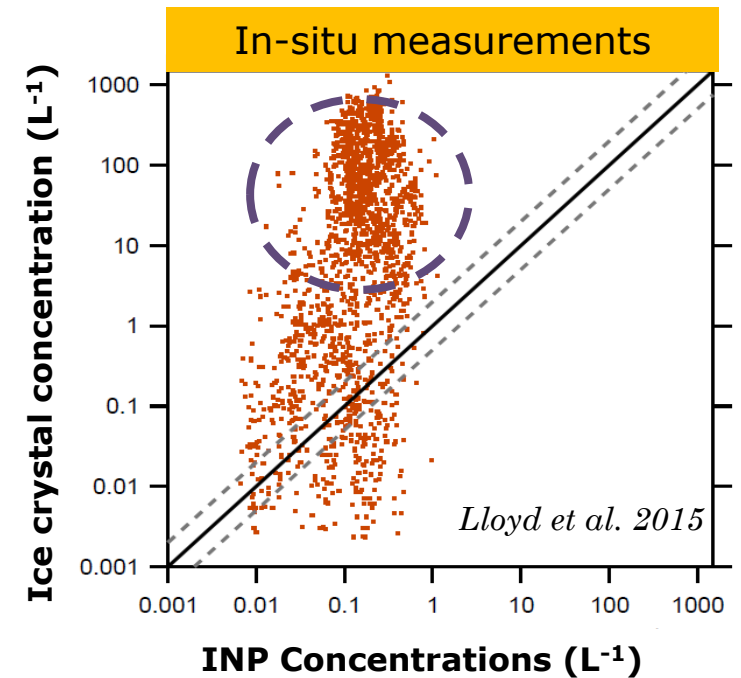


NW winds

- ✓ Very steep ascent
- ✓ Stronger updrafts
- ✓ High supersaturations
- ✓ Frequent mixed-phase conditions



Blowing Snow: surface winds lift snowflakes which can provide ice crystals to the clouds above



During the CLACE2014 campaign at Jungfraujoch (Alps), **blowing snow** was proposed to be responsible for the enhanced ice number concentrations when:

- ✓ Wind speed $\gtrsim 5 \text{ ms}^{-1}$
- ✓ Ice number concentrations $\lesssim 100 \text{ L}^{-1}$

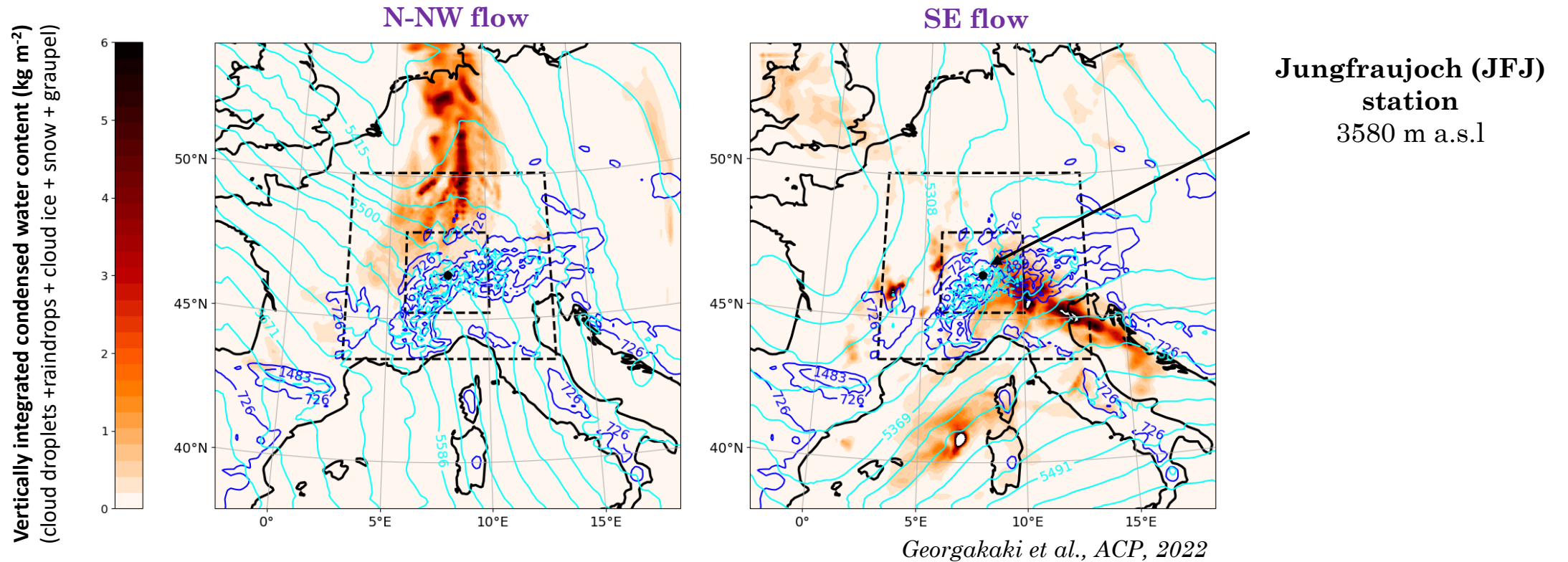


SIP processes in orographic MPCs Modeling with WRF

CLACE 2014 field campaign

“Cloud Aerosol Characterization Experiments” (Lloyd et al. 2015)

- ✓ 3 nested domains surrounding JFJ with horizontal resolution: **12km - 3km - 1 km**



Two simulation periods in order to investigate the dynamical influence caused by the local orography:

- ✓ 25.01.2014 00.00 UTC – 28.01.2014 00.00 UTC : steep ascent of the airmasses before arriving at JFJ
- ✓ 29.01.2014 00.00 UTC – 01.02.2014 00.00 UTC : gentle ascent of the airmasses over the Aletsch Glacier

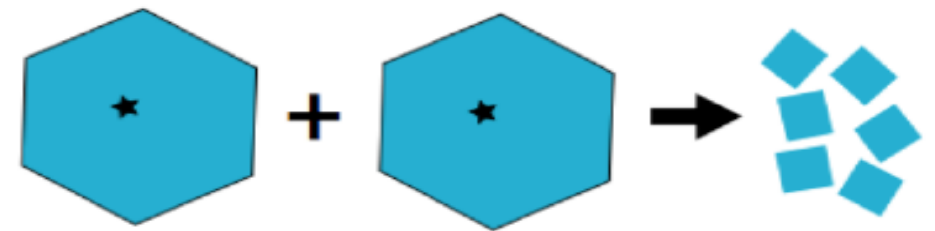
Collisional Break-up

Morrison scheme: 2-moment bulk microphysics, 5 hydrometeor species:



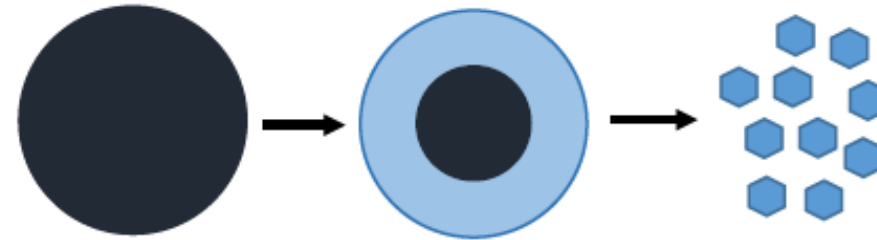
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- (5) graupel – graupel collisions → fragmentation of graupel



Fragments are added to **cloud ice** category

Freezing and shattering droplets



Following Phillips et al. (2018)

Fragmentation of freezing raindrops is assumed to occur during:

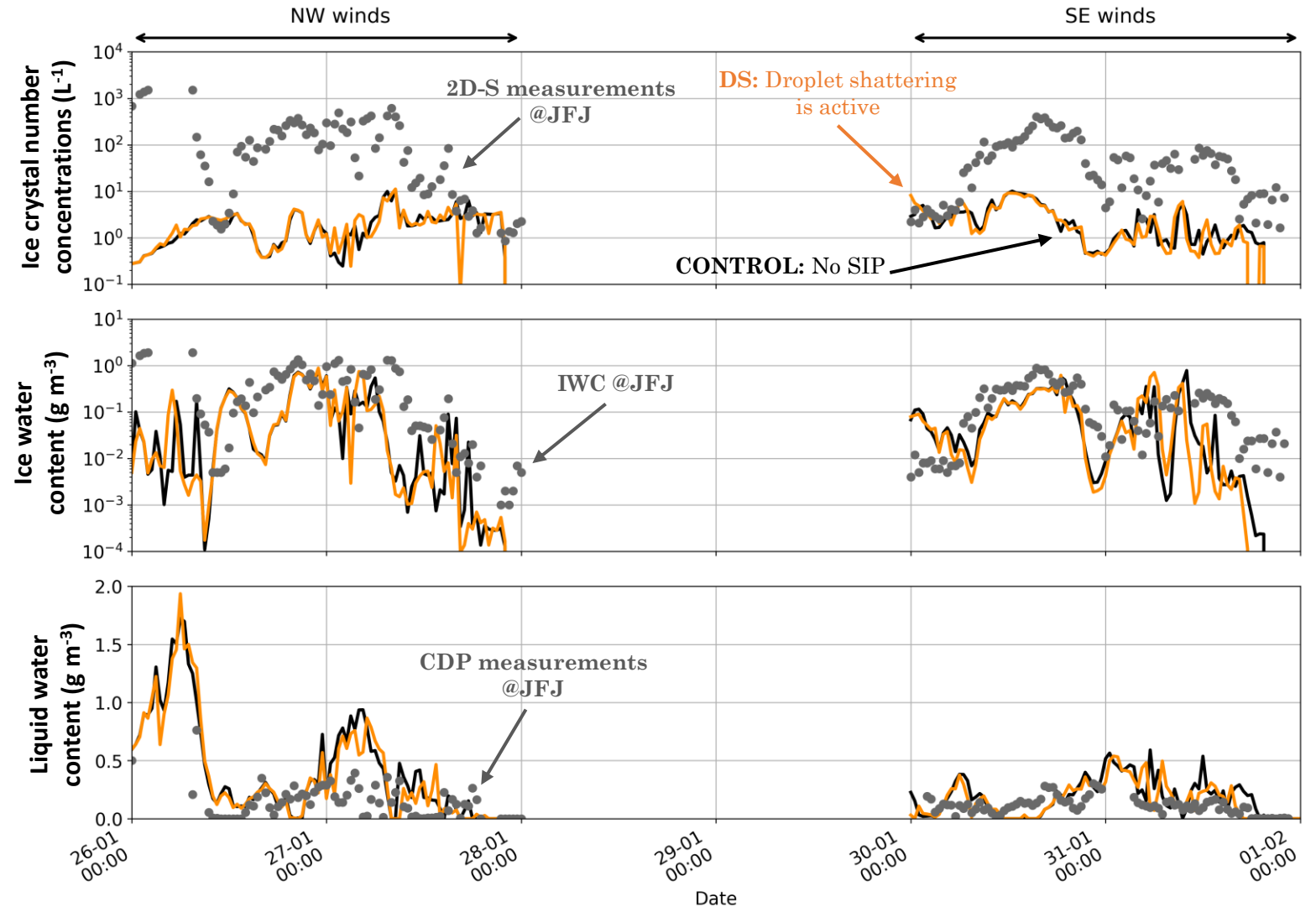
- MODE 1:** Collision with smaller ice-particle (cloud ice) or heterogeneous drop freezing → both **BIG** and **tiny** fragments will be emitted as a function of **temperature** and **droplet size** → tiny ejected splinters initiated as new cloud ice
- MODE 2:** Collision with more massive ice-particle (snow/graupel) → only tiny splinters are generated as a function of **temperature**, **droplet size** and **collisional kinetic energy**.

$$N_{DS} = \mathbb{E}(\mathbf{D})\Omega(\mathbf{T}) \left[\frac{\zeta\eta^2}{(\mathbf{T} - \mathbf{T}_0)^2} + \beta\mathbf{T} \right]$$

$$N_{DS} = 3 \Phi(\mathbf{T}) \times [1 - f(\mathbf{T})] \times \max(\mathbf{DE} - DE_{crit}, 0)$$

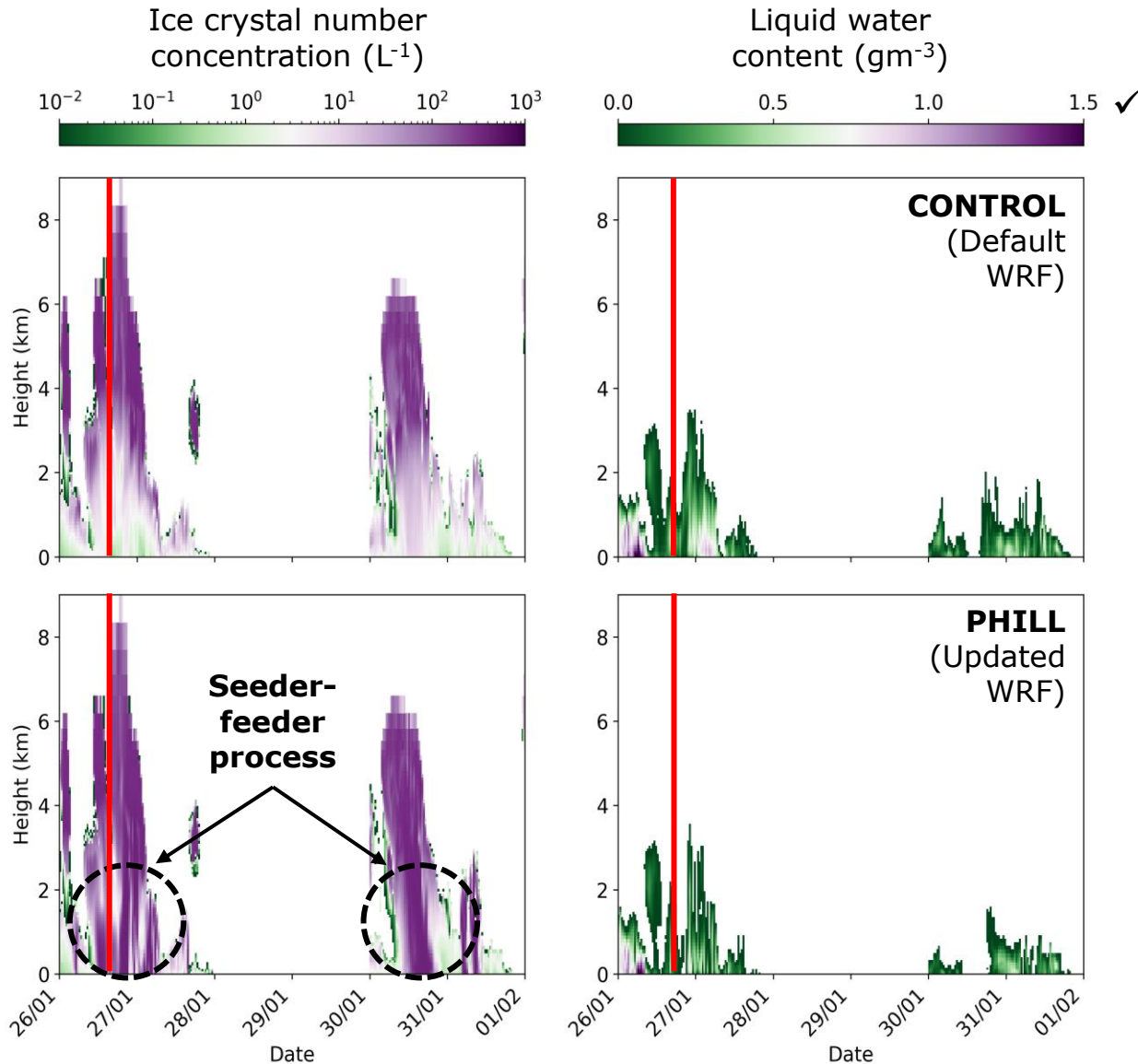
CLACE 2014: effect of DS mechanism alone

- ✓ CONTROL (black line) and DS (orange line) underestimate the ice crystal number concentrations by 2-3 orders of magnitude
- ✓ The modeled ice water content is outside the observed range
- ✓ The cloud liquid water is overestimated



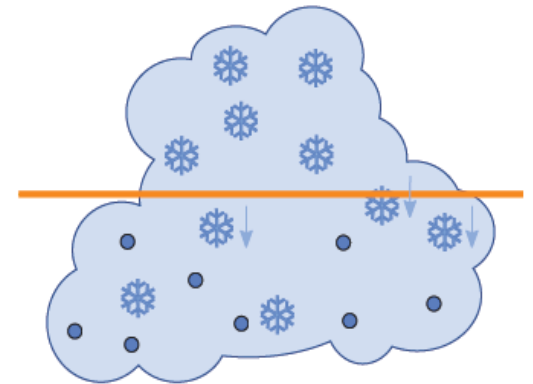
CLACE 2014: effects from ice-ice collisions

External seeder-feeder process



Seeding ice crystals can enhance the collision efficiencies and hence SIP through BR in the lower-lying parts of the cloud

In-cloud seeding



Proske et al. 2021

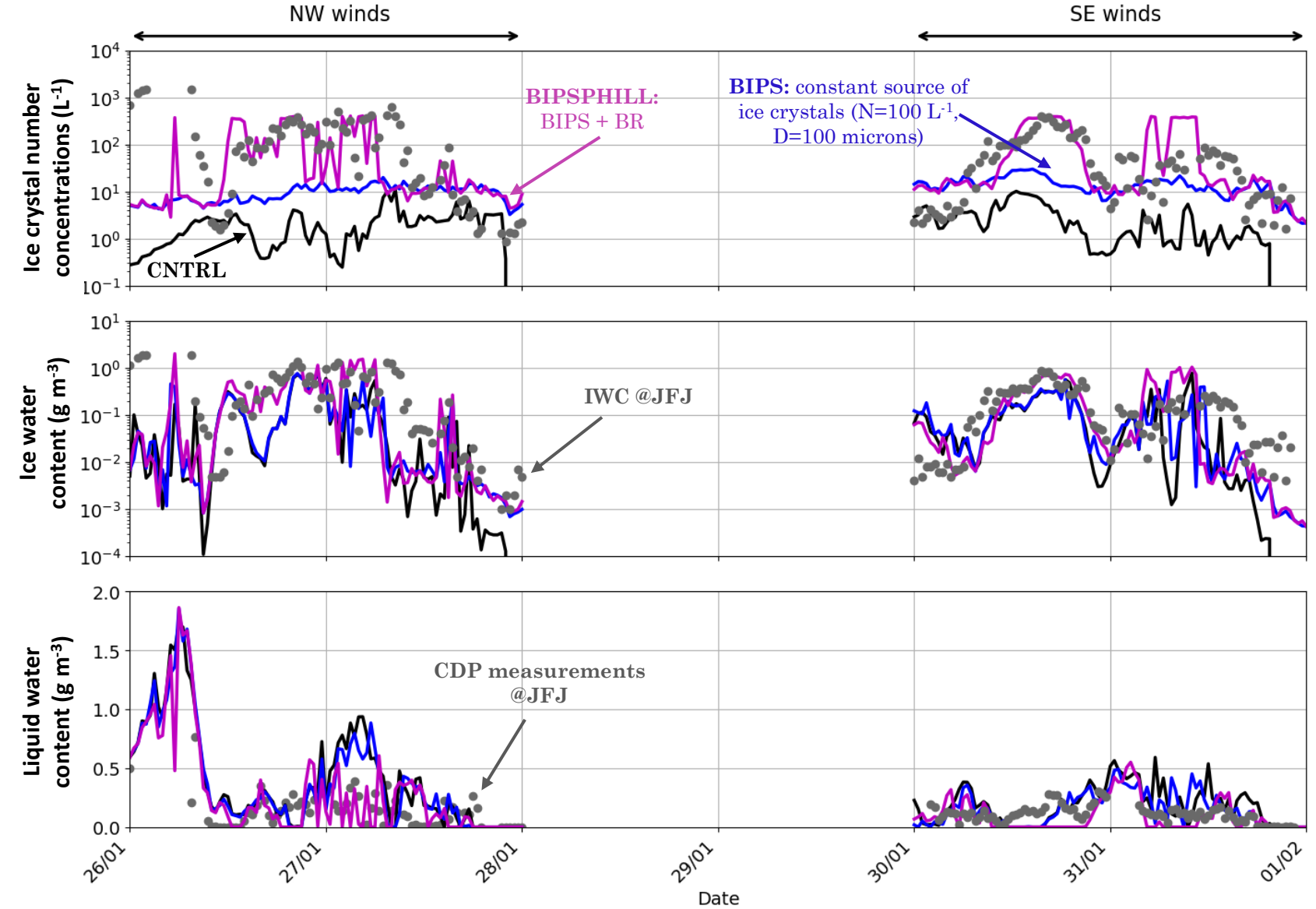
✓ Satellite observations reveal that **seeder-feeder situations** (both external and in-cloud) occur frequently over Switzerland (Proske et al., 2021)

CLACE 2014: impact of “blowing snow”

Blowing snow representation:

Constant source of ice crystals (100 L^{-1}) with sizes of 100 microns in the first cell of WRF

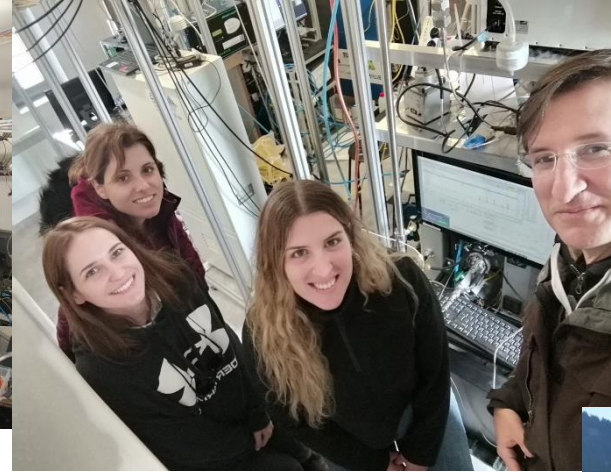
- ✓ Blowing snow alone (light blue line) cannot account for the observed ice particle concentrations, as it never predicts ICNCs $\geq 50 \text{ L}^{-1}$
- ✓ “Blowing snow” is important when a cloud is near the ground (i.e., ice supersaturated environment)
- ✓ The combined effect of blowing snow and collisional break-up (magenta line) results in best agreement with measured ICNCs



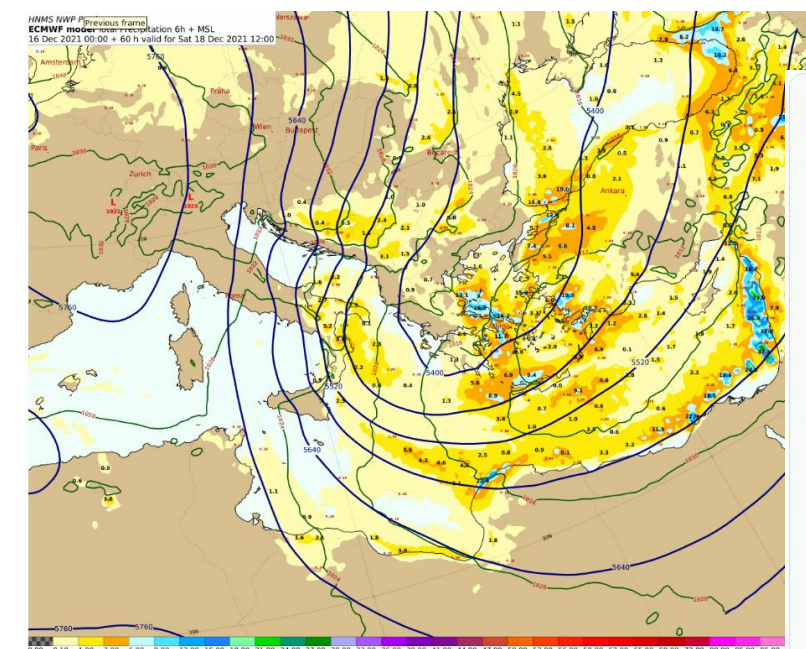
CALISHTO campaign: <https://calishto.panacea-ri.gr/>



Cloud-Aerosol InteractionS in the Helmos background Troposphere (Oct.21-Feb.22)



Case study: Storm Carmel visiting Greece

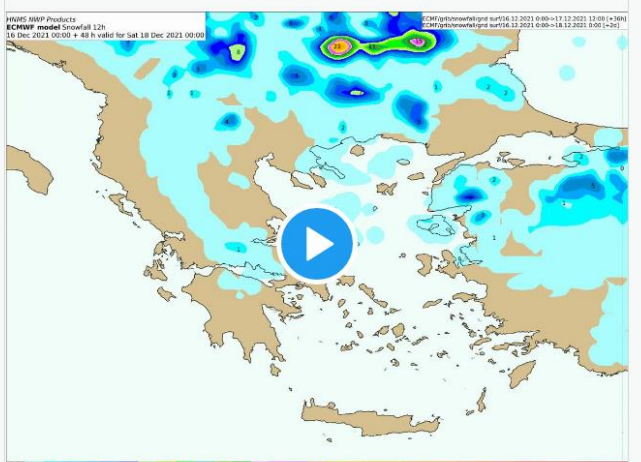


Keep Talking Greece
@keptalkingGR · Follow

#Carmel it is

Theodoros Kolydas @KolydasT

Σε λίγο θα εκδοθεί #έκτακτο δελτίο #επιδείνωσης .Το Τμήμα Μετεωρολογίας της Κύπρου σε συνεννόηση με #ΕΜΥ και Μετεωρολογική Υπηρεσία του #Ισραήλ αποφάσισε όπως δώσει το όνομα #Carmel στην συγκεκριμένη κακοκαιρία. Στο βίντεο οι περιοχές που θα επηρεαστούν από χιονοπτώσεις @GSCP_GR



3:11 PM · Dec 16, 2021



Snowfall Forecast by the National Meteorological Service (EMY)

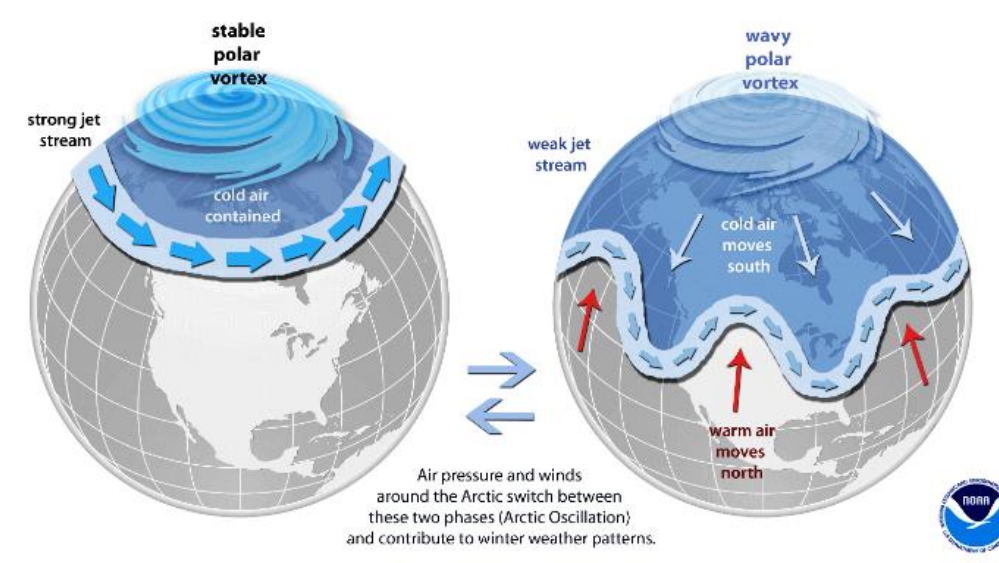
WEATHER WARNING: CARMEL TO HIT GREECE WITH SNOW, RAIN, STORMS, WIND 9 B

December 16, 2021 weather Comments Off

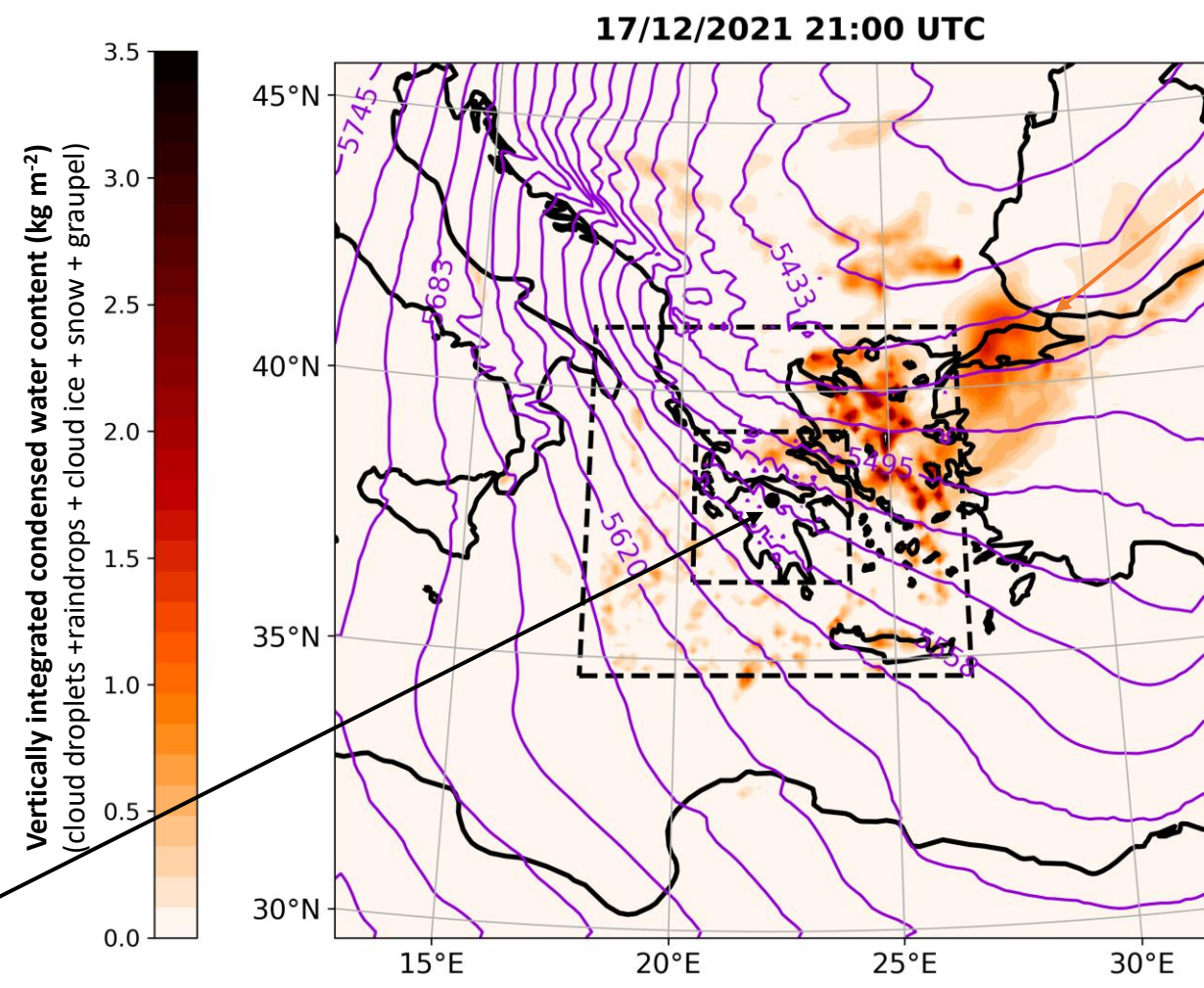
Sharp temperature drop, stormy winds and snowfall in the central and southern parts of Greece on **December 18th**.

The “culprit” behind the storm?

...a visit from the polar vortex breaking out of the Arctic



Modeling Storm Carmel with the Weather Research and Forecasting model (WRF)

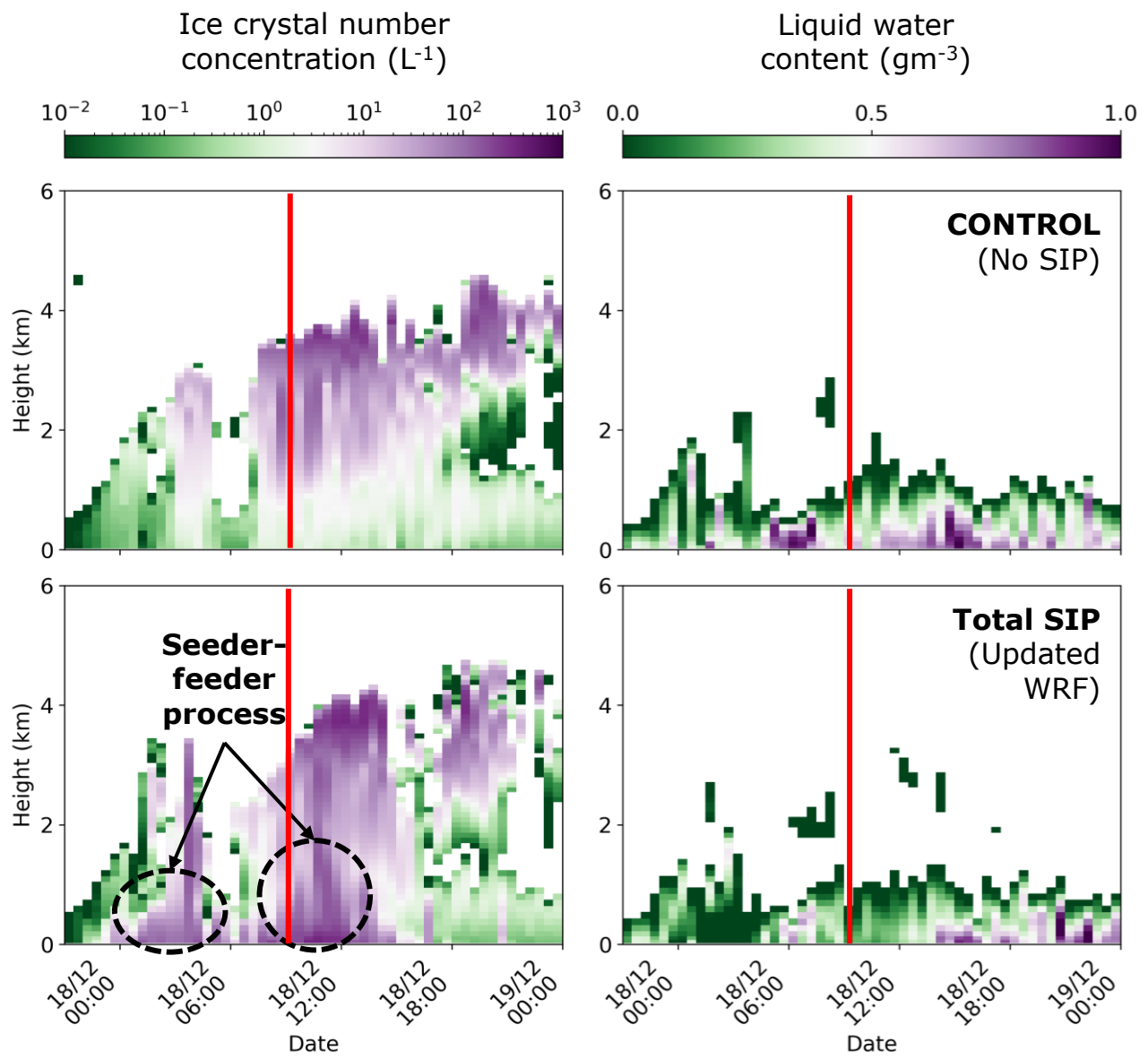


Polar airmasses arriving from the Northeast

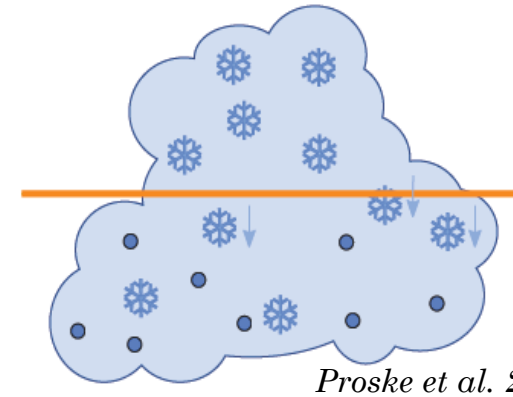
reference lat/lon = 37.984°N, 22.196°E (@HAC station)

- 3 nested domains surrounding the HAC station with horizontal resolution: **12km - 3km - 1 km**
- **Simulation period:** 17.12.2021 00.00 UTC – 19.12.2021 00.00 UTC
- Spin-up time: 21 hours
- **Time-step:** 36s-9s-3s

Conditions favoring ice multiplication at Mt Helmos



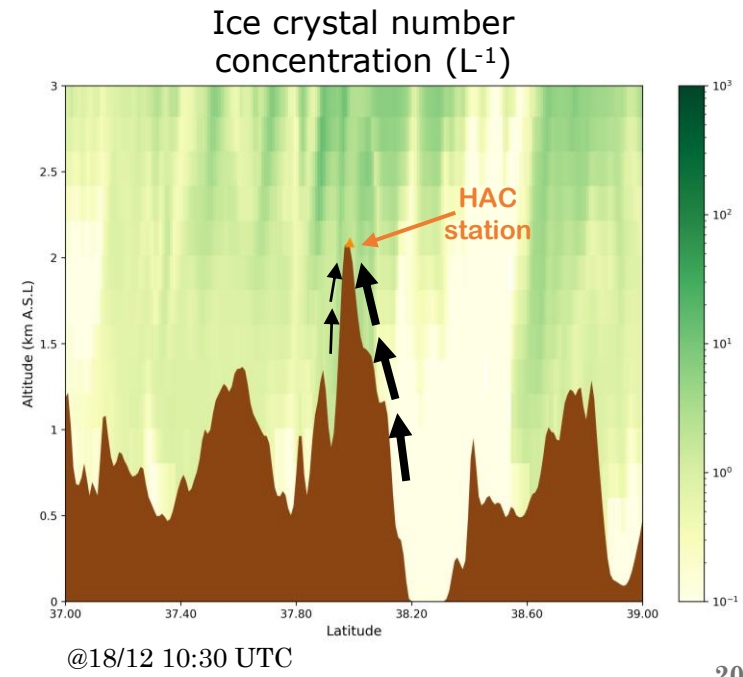
In-cloud seeding



Proske et al. 2021

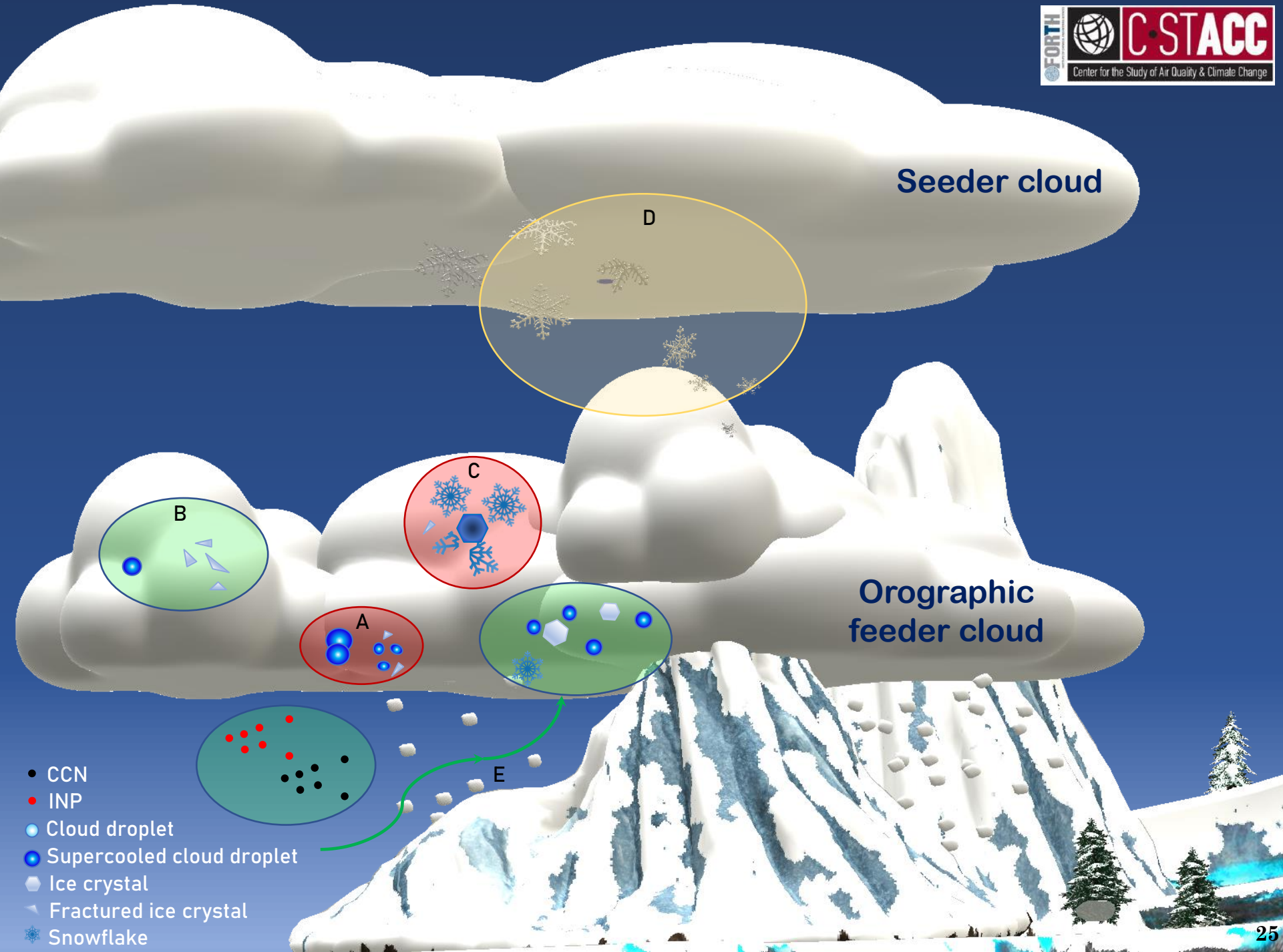
1. Seeding ice crystals enhance the collision efficiencies and SIP through BR in the lower-lying parts of the cloud

2. Steep ascent of the airmasses due to strong orographic forcing → mixed-phase conditions maintained

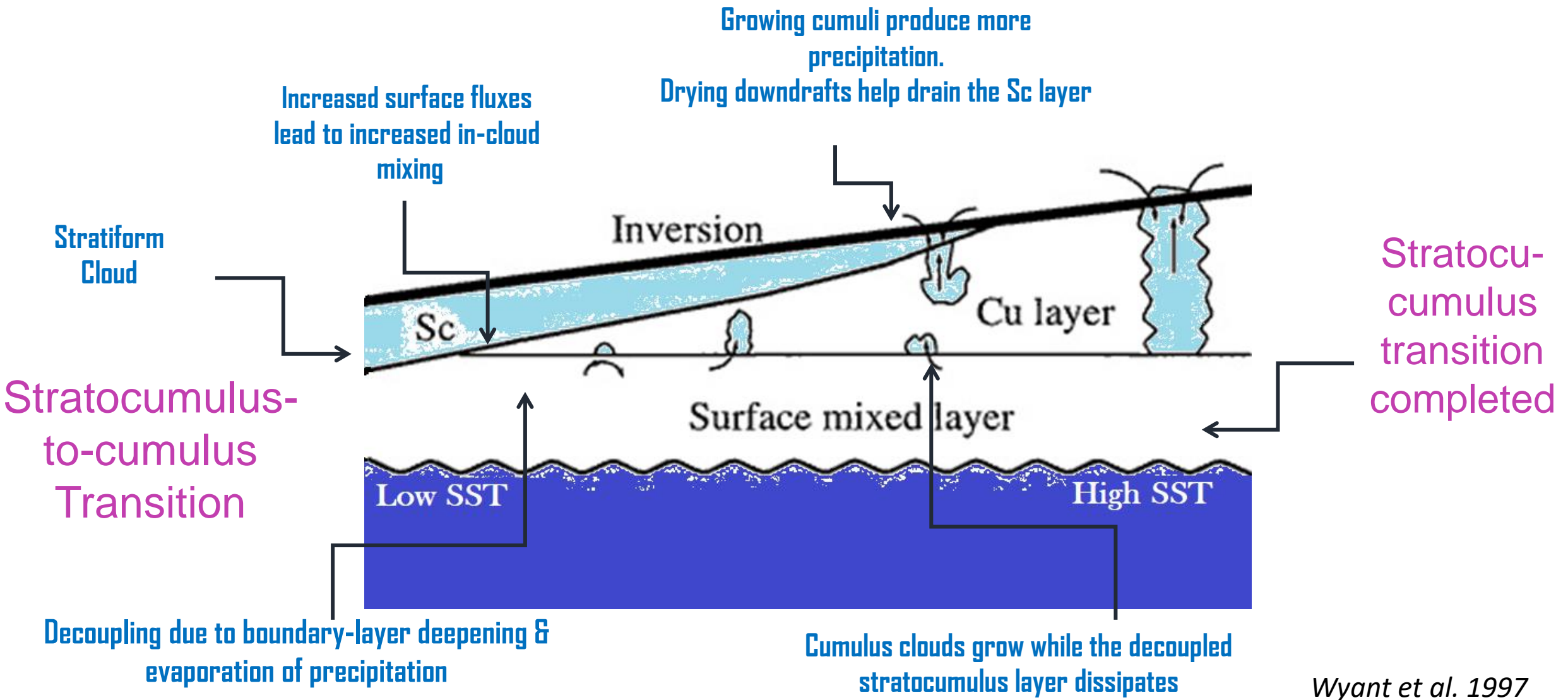


Wintertime orographic mixed-phase clouds

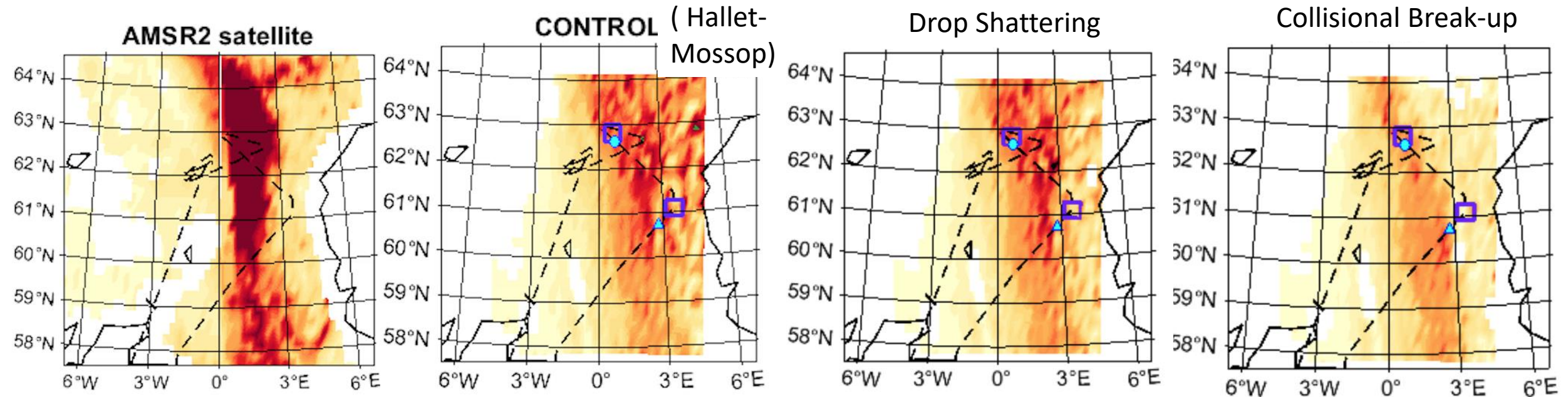
- ✓ The Hallett-Mossop process → ineffective since $T < -8\text{ }^{\circ}\text{C}$ [A]
- ✓ The Droplet Shattering process → not efficient due to a lack of big raindrops [B]
- ✓ The collisional break-up process → elevates ICNCs up to 3 orders of magnitude but is activated in certain cases [C]
- ✓ The seeder feeder effect → frequent over Switzerland → enhanced collision efficiency in the low-level feeder clouds → further promotes cloud glaciation [D]
- ✓ Blowing snow → significant contribution when cloud is near the ground → further facilitates SIP through BR [E]



SIP effects on Stratocumulus- to-Cumulus Transition (Cold air outbreaks)



Simulations of a CAO event (with WRF) observed north of UK (Nov.2013)

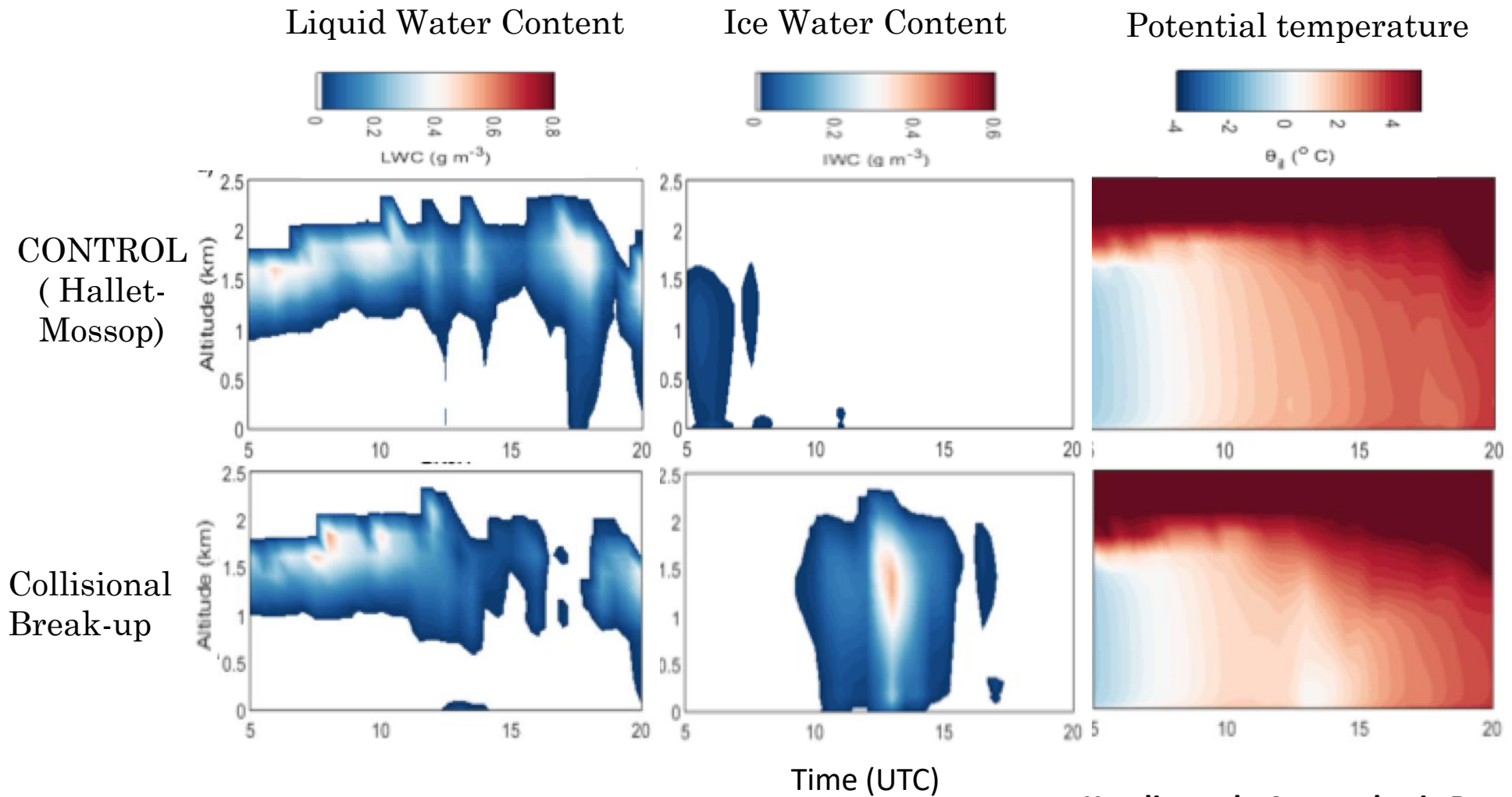


Liquid Water Path

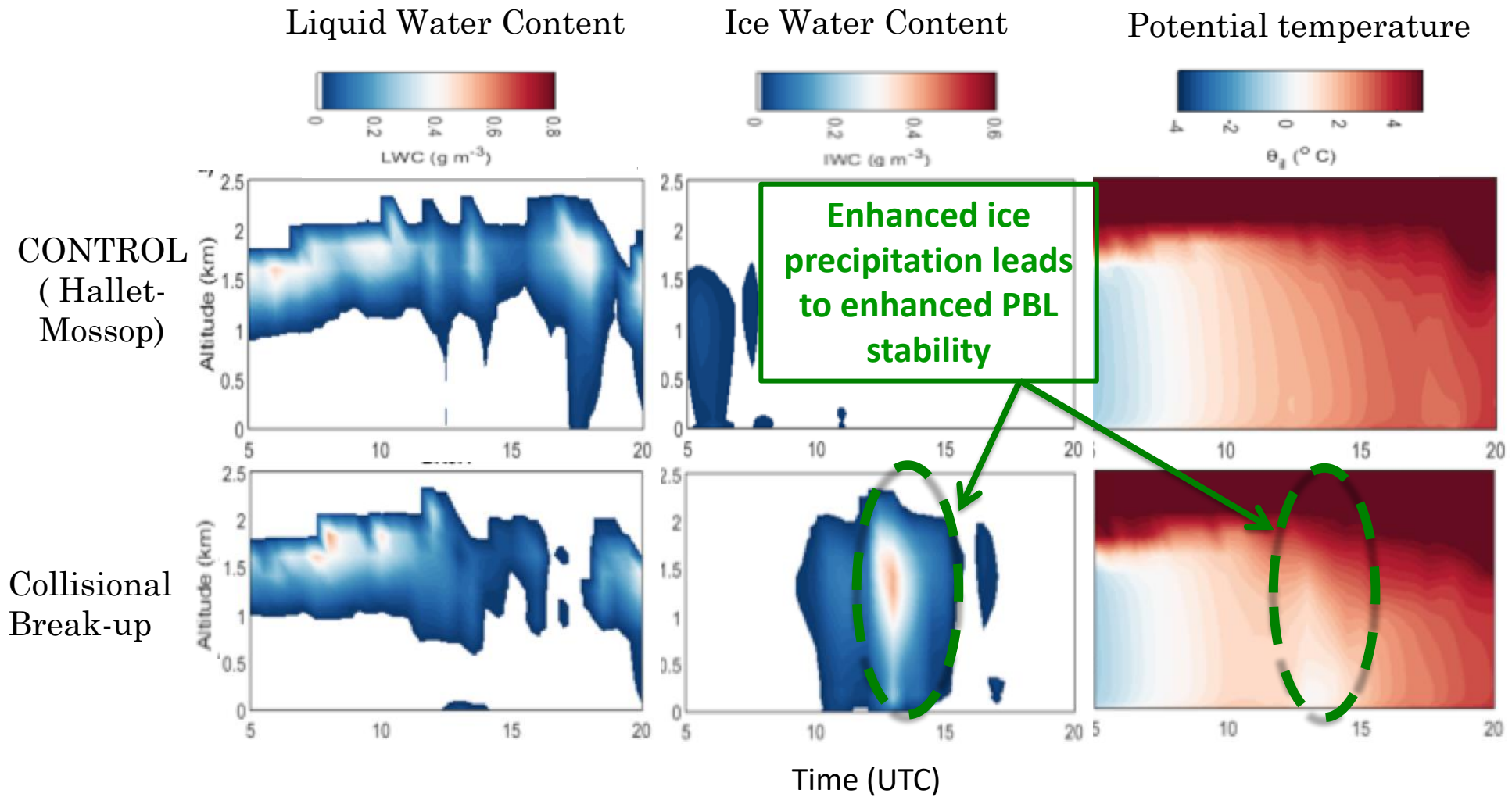
Karalis et al., Atmospheric Research (2022)

- ✓ Standard WRF (CONTROL) cannot reproduce the Stratocumulus-to-Cumulus transition correctly
- ✓ Increasing ice production through drop-shattering & especially collisional break-up improves the transition

SIP effects on Stratocumulus- to-Cumulus Transition (Cold air outbreaks)



SIP effects on Stratocumulus- to-Cumulus Transition (Cold air outbreaks)



Activation of the missing SIP mechanisms (*collisional break-up mostly*)



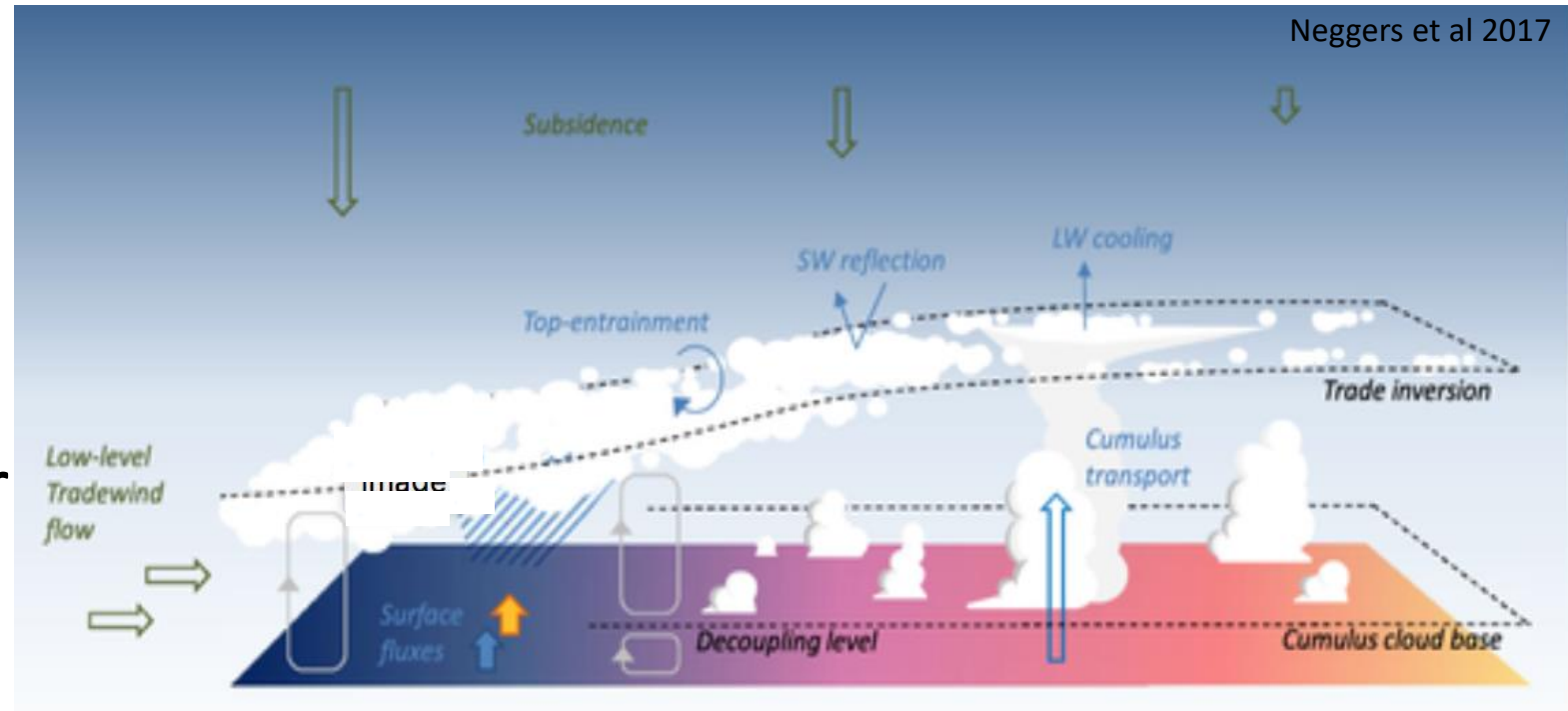
Enhanced precipitation



Enhanced boundary-layer decoupling

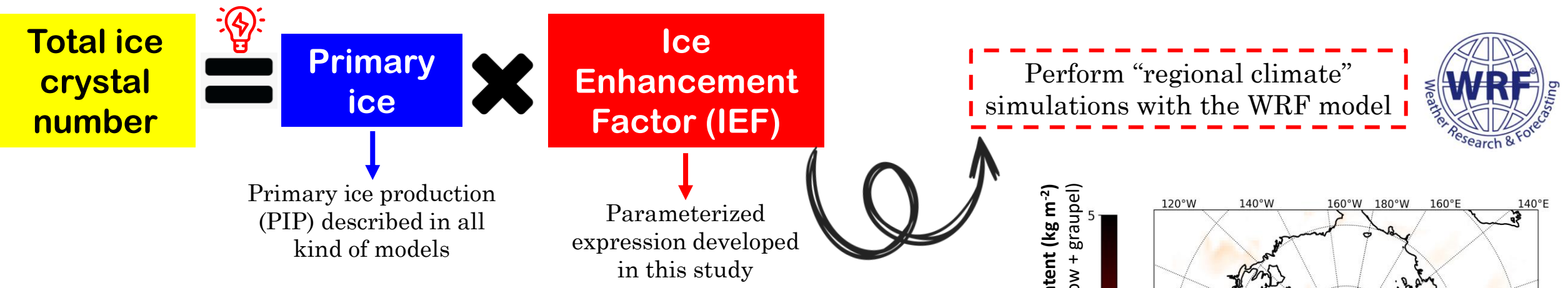


Accelerated onset of the stratocumulus-to-cumulus transition

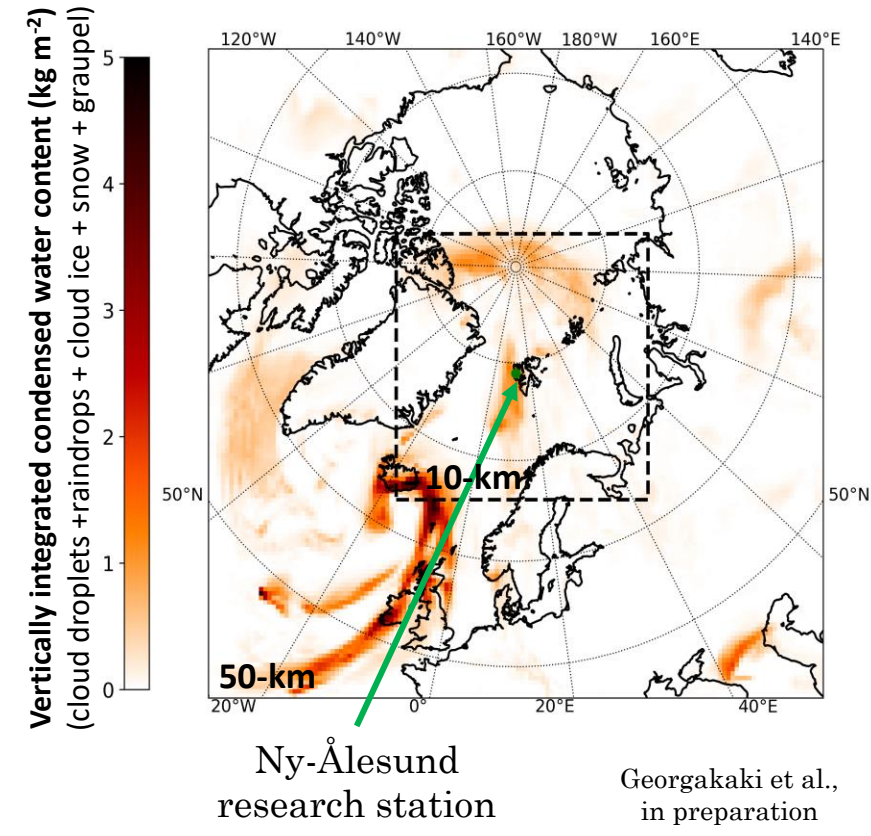


Karalis et al., Atmospheric Research (2022)

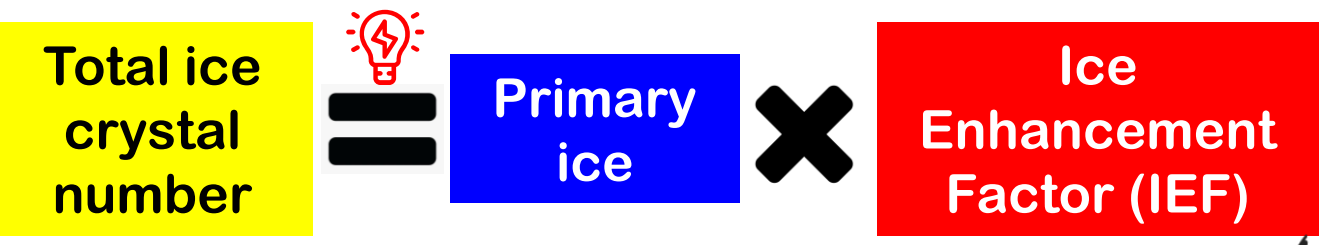
Our approach to parameterize SIP in polar stratiform clouds



- ✓ Focus on Arctic mixed-phase clouds
- ✓ 2-year simulation period: 2016-2017, Ny-Ålesund
- ✓ Ice processes in the updated version of WRF:
 - i. Primary ice production
 - Homogeneous freezing
 - Heterogeneous freezing
 - ii. Secondary ice production
 - Hallett-Mossop
 - Collisional break-up
 - Droplet shattering



Our approach to parameterize SIP in polar stratiform clouds



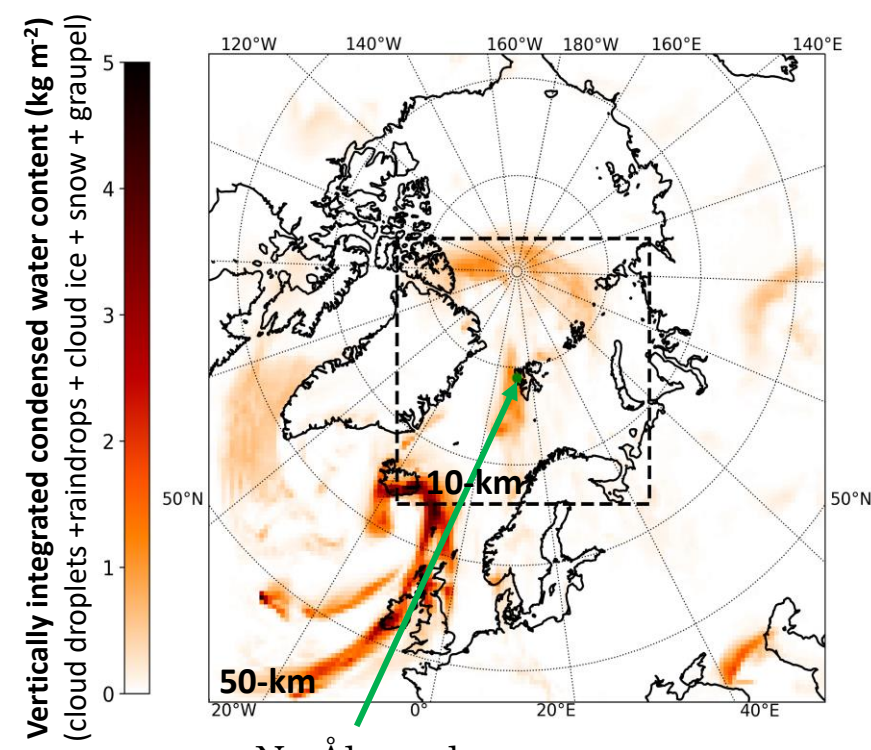
Perform “regional climate” simulations with the WRF model



- ✓ Outputs extracted from the 10 km-resolution (nest)
- ✓ IEF encompasses the effect of all 3 important SIP processes:

$$IEF = IEF_{BR} + IEF_{DS} + IEF_{HM}$$

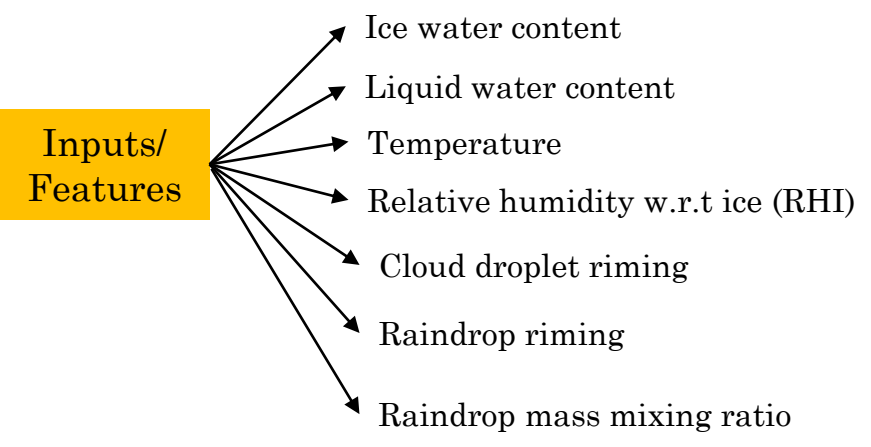
$$IEF_i = 1 + \frac{SIP_i \text{ rate}}{PIP \text{ rate}}$$



Ny-Ålesund research station

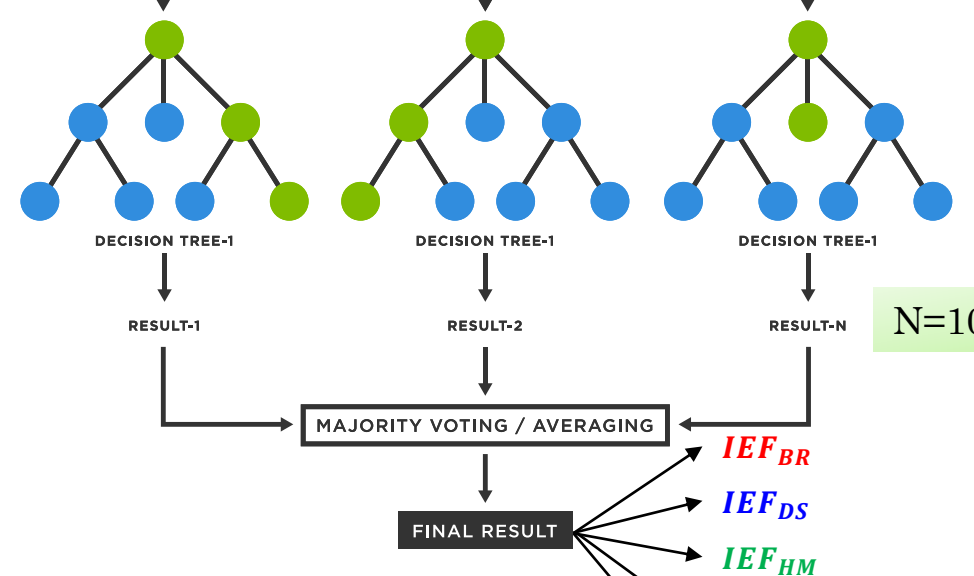
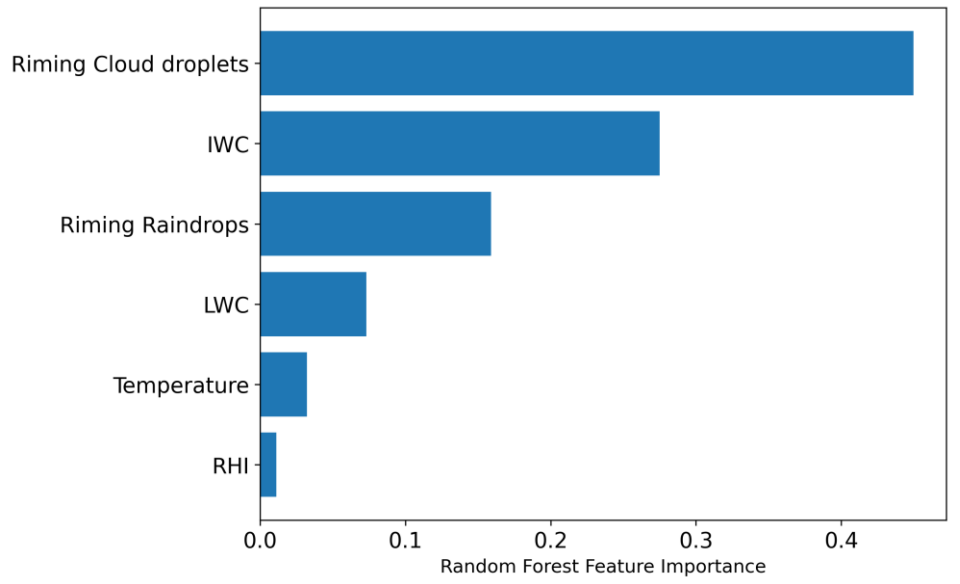
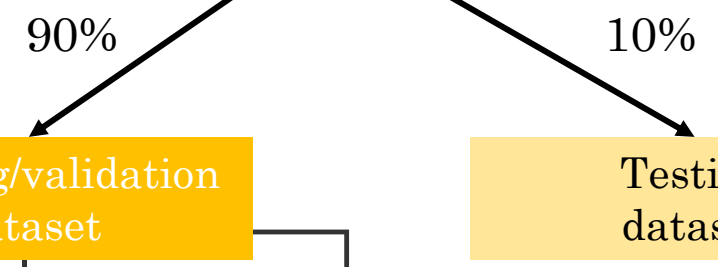
Georgakaki et al., in preparation

Developing SIP parameterization: Random Forest (RaFSIP) regressor



Developed using the RandomForestRegressor class from the **scikit-learn** package

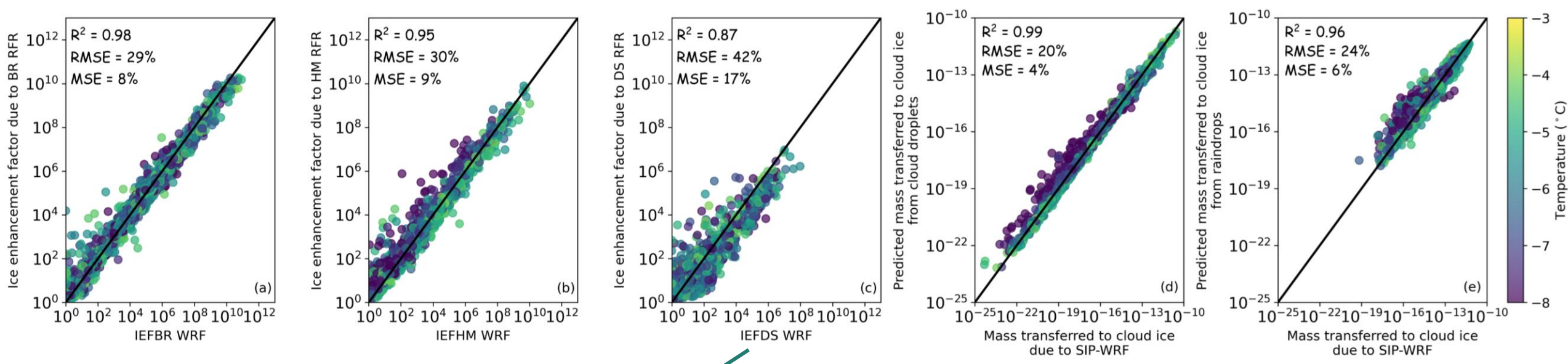
WRF 2-year high-resolution outputs



- IEF_{BR}
- IEF_{DS}
- IEF_{HM}
- Mass transferred from cloud droplets to cloud ice
- Mass transferred from raindrops to cloud ice

Offline performance of the RaFSIP algorithms

Conditions		RaFSIP model	RaFSIP predictions			
Temperature	Rain mass mixing ratio		IEFBR	IEFHM	IEFDS	
$-8 \leq T < -3^\circ\text{C}$	> 0	forestALL	✓	✓	✓	✓

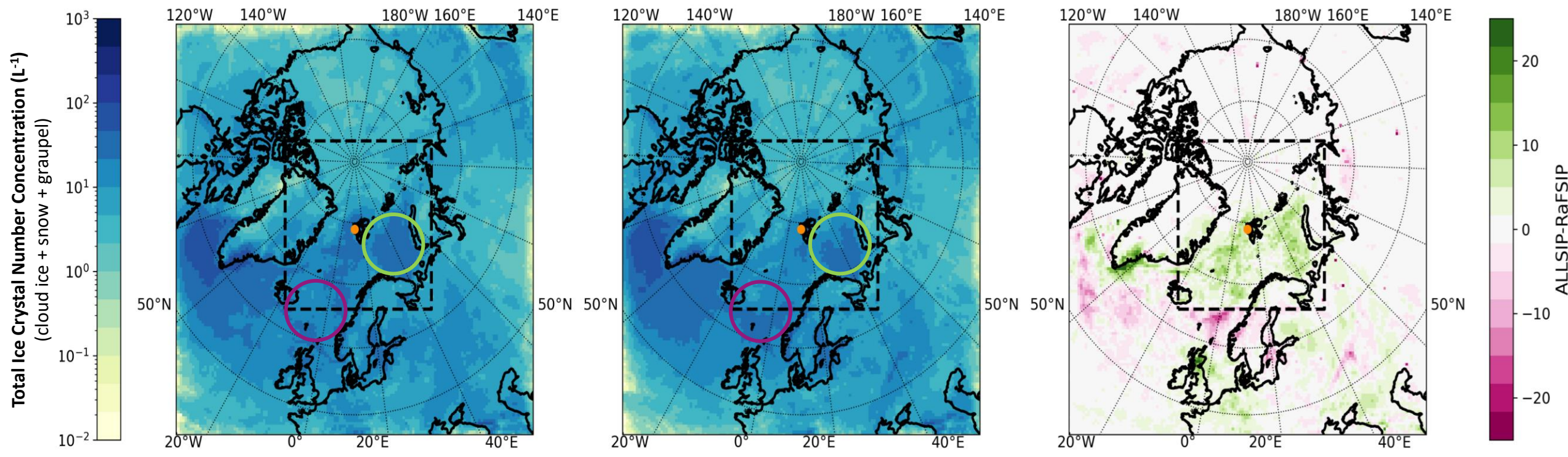


Less constrained IEF due to DS probably because of less training examples compared with the other SIP processes

Online performance of the new RaFSIP parameterization

WRF simulation with detailed inline microphysics

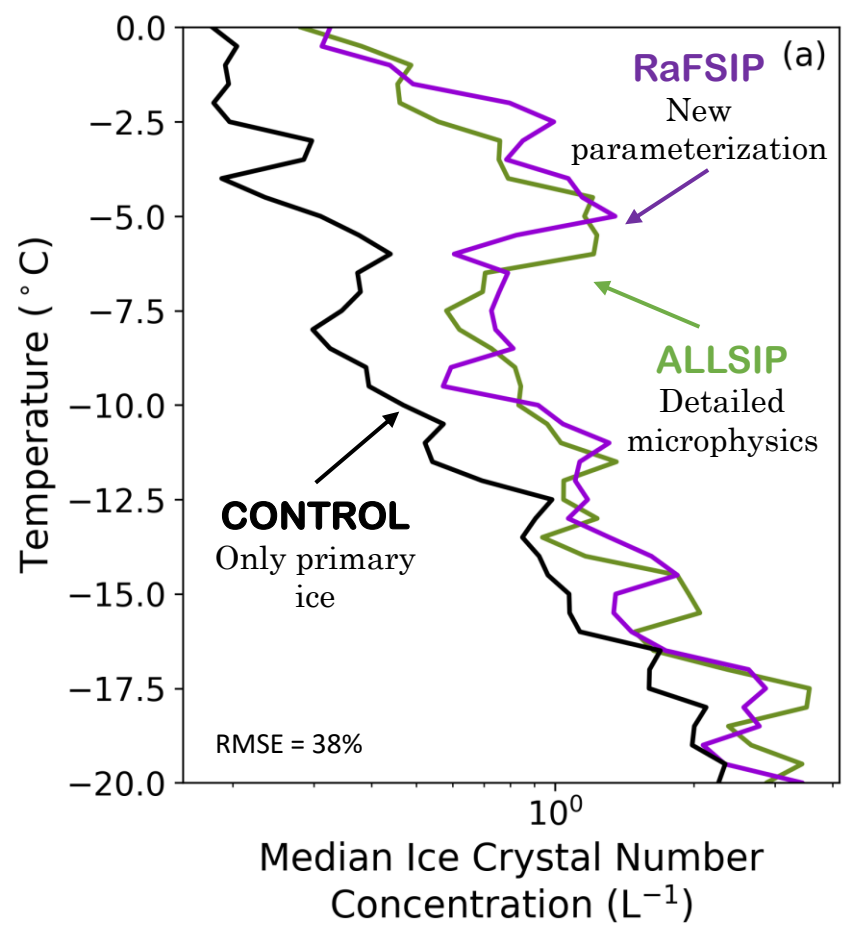
WRF simulation with RaFSIP parameterization



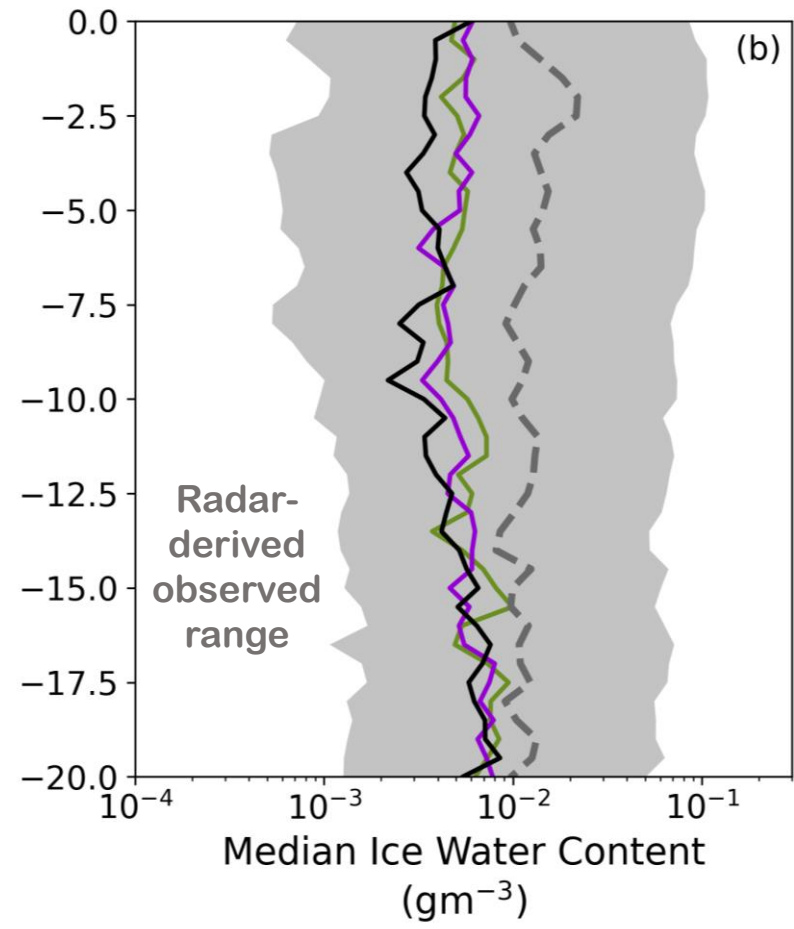
* Averaged concentrations within the PBL

- ✓ Good performance over the continental regions
- ✓ Certain regions where the new RaFSIP leads to **overestimations/underestimations** mainly over sea-ice → the RaFSIP is trained over continental grid points

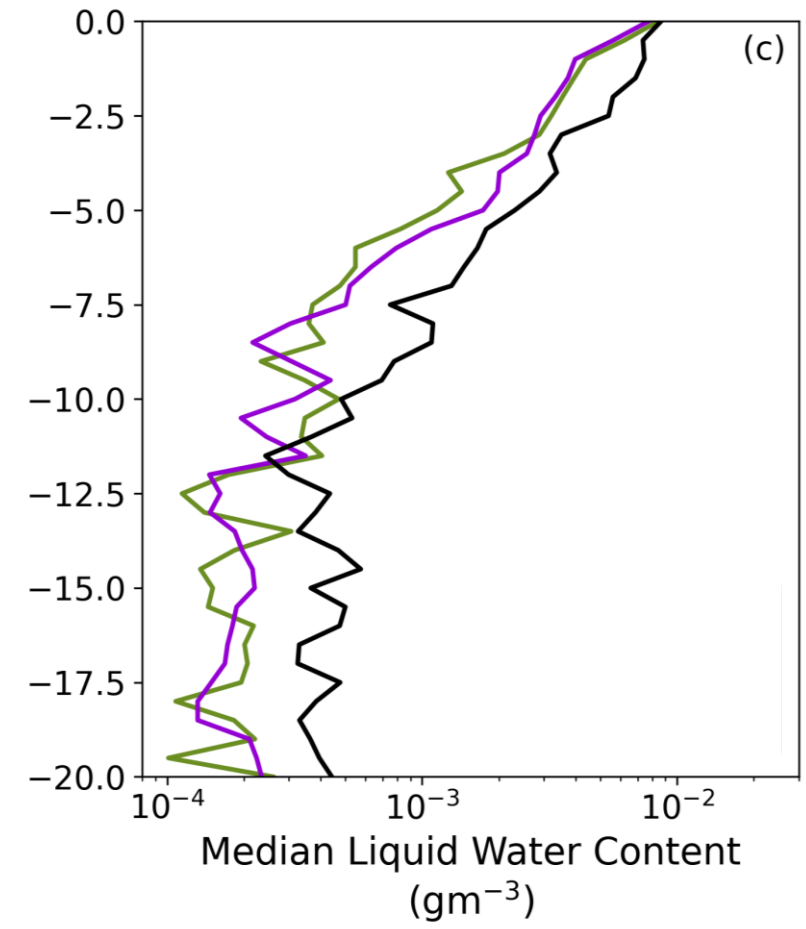
Focusing on the Ny-Ålesund model grid point



✓ A factor of up to ~5 increase in the ice crystal concentrations



✓ Not significant change in IWC due to a shift towards smaller ice particles

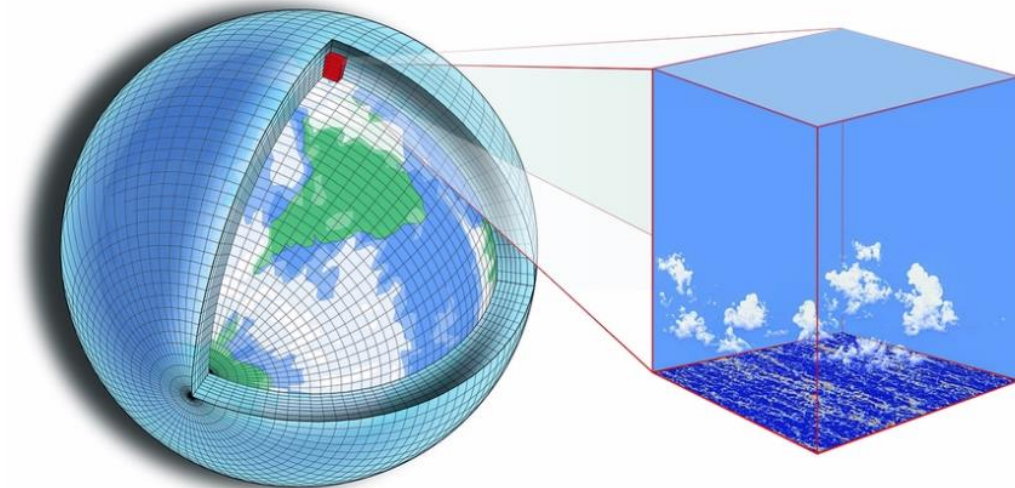


✓ Slight ↓ in LWC through the ice growth properties (WBF + riming) of the small secondary ice particles

FORCeS Ice Experiment (FOR-ICE)

Climate model intercomparison project (NorESM2, ICON-HAM, EC-Earth) to quantify the sensitivity to:

- ✓ Ice Nucleation
- ✓ Secondary ice production
- ✓ Sedimentation
- ✓ Wegener-Bergeron-Findeisen process
- ✓ Large-scale, convective & turbulent ice transport

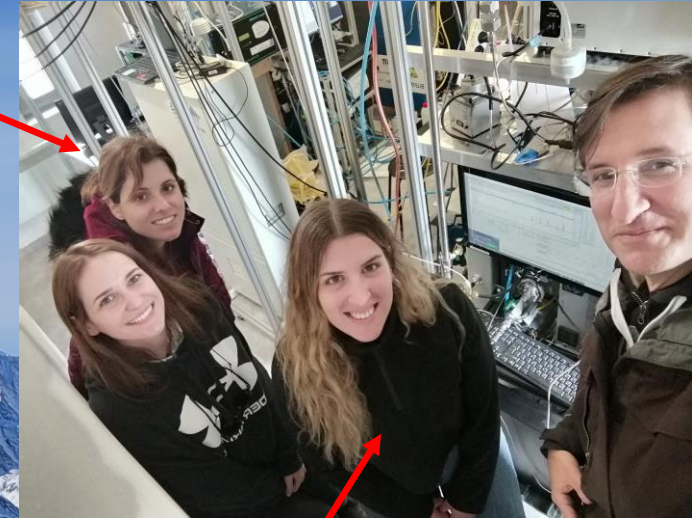


- Primary Ice Production is critical to “get right”, no matter what.
- Secondary Ice Production can have larger influence on ice formation than primary ice at $T < -15^{\circ}\text{C}$. *It seems to be acting everywhere we looked at.*
- In Polar and Orographic environments, “seeder-feeder” configurations can lead to considerable secondary ice production. *Completely new view...*
- Enhanced precipitation rates associated with secondary ice can affect the development of larger-scale cloud systems, such as Statocu-to-Cu transitions during cold air outbreaks.
- Secondary ice processes are highly uncertain, but can affect most types of MPC with important implications for radiation, precip and glaciation fraction.



Prof. Sylvia Sullivan

Dr. Georgia Sotiropoulou



Paraskevi Georgakaki

THANK YOU!

