Overview of our atmospheric INP measurements from polar to tropical regions and what we might learn from them.

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and many colleagues from TROPOS and cooperating research groups

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Climate Change Canada

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overview

- (very) short introduction on INP in general

- a bit about INP measurements & data evaluation
- where / what we measured, wrt. INP in
 - Beijing
 - the Arctic
 - Cabo Verde

- what I think might be true about INP in general



ice nucleating particles (INP)

a first review on materials that may occur in the atmosphere and have shown ice nucleating activity:

Szyrmer, W., and I. Zawadzki (1997), Biogenic and anthropogenic sources of ice-forming nuclei: A review, *BAMS*, 78(2), 209-228.



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more recent reviews:

- Murray et al. (2012), Ice nucleation by particles immersed in supercooled cloud droplets, *Chem. Soc. Rev.*

- Hoose, C., and O. Möhler (2012), Heterogeneous ice nucleation on atmospheric aerosols: A review of results from laboratory experiments, *Atmos. Chem. Phys.*

- Kanji et al. (2017), Chapter 1: Overview of Ice Nucleating Particles, in *Ice Formation and Evolution in Clouds and Precipitation: Measurement and Modeling Challenges, edited, Meteor. Monogr.*

- Coluzza et al. (2017), Perspectives on the future of ice nucleation research: Research needs and unanswered questions identified from two international workshops, *Atmosphere*



ice nucleating particles (INP)

some basics known today (some of them corroboating what was formerly known already):

- mineral dust particles are important atmopsheric INP
- among mineral dusts, K-feldspar particles may play the most important role
- biogenic INP are macromolecules (proteins or polysaccarides) found on microorganisms
- sea spray may contribute to atmospheric INP, but to what extent is still open



concentrations of ice nucleating particles (N_{INP})



Petters & Wright (2015), Revisiting ice nucleation from precipitation samples, GRL.



measuring ice nucleating particles (N_{INP})

<u>in-situ</u>

- continuous flow diffusion chambers (CFDC, with names like PINC, SPIN, HINC, INKA, ...)
- laminar diffusion tube (LACIS)
- expansion chambers (e.g., AIDA, CLOUD-Chamber, PINE)

- characteristic: optical detection of the ice crystals; due to this and the limited flow, comparably high INP number concentrations have to be present ($\sim > 1/L$); often inlet cut-off of a few micrometers



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<u>off-line</u>

- collecting particles on a filter (or using a batch sample)

- analyzing them off-line in "cold-stages" or "freezing arrays"; different types are now operated by a growing number of different groups

- characteristic:
 - + comparably low concentrations can be detected
 - contaminations limit the lowest temperature to which measurements can be made (depending on droplet size and other factors, > -25°C; exception: pico-liter droplets, but droplet production and optical detection much more expensive)
 - long sampling times -> bad time resolution



measuring ice nucleating particles (N_{INP})



measurement principles of our cold-stage / freezing array:

INDA and LINA are used for suspensions (e.g., washed polycarbonate filters),

INDA also for filter punches (quartz-fiber filters),

LINA reaches lower temperatures





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* Leipzig Ice Nucleation Array ****** Ice Nucleation Droplet Array









Vali, J. Atmos. Sci. (1971)

*f*_{ice} -> fraction of frozen droplets from all droplets
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cumulative distribution 1 -0.1 $f_{\rm ice}$ 0.01 -15 -25 -20 -10 -5 T[°C]



 $f_{\rm ice}$ -> fraction of frozen droplets from all droplets



Vali, J. Atmos. Sci. (1971)





λ

Vali, J. Atmos. Sci. (1971)

- -> for us: average number of ice active entities (INP) per droplet
- f_{ice} -> fraction of frozen droplets from all droplets



cumulative distribution 1 -**Poisson distribution:** 10⁻² 0.1 $\lambda = -\ln(1 - f_{ice})$ [1/L] ^{dNI}N $f_{\rm ice}$ (e.g., $f_{\rm ice} = 0.4 \rightarrow \lambda = 0.51$) 10⁻⁴ 0.01 Normalization: -20 -25 -15 -10 -5 -25 -15 -20 -10 -5 T[°C] *T* [°C] $N_{\rm INP} = \lambda / V_{\rm droplet} = - \ln (1 - f_{\rm ice}) / V_{\rm droplet}$ Vali, J. Atmos. Sci. (1971) λ -> for us: average number of ice active entities (INP) per droplet -> fraction of frozen droplets from all droplets $f_{\rm ice}$ TROPOS 9 Heike Wex, 4.8.2020 @ online INP colloquium

Polen et al. (2018), Cleaning up our water: ..., AMT.

see supplement of Wex et al. (2019), ACP.

 $N_{INP,corr} = (-ln(1 - f_{ice,s}) + ln(1 - f_{ice,b}))/V$

Poisson distribution: 10⁻² $\lambda = -\ln(1 - f_{ice})$ [7/1] ^{dNI}N $(e.g., f_{ice} = 0.4 \rightarrow \lambda = 0.51)$ 10⁻⁴ Normalization: -20 -10 -5 -25 -15 -10 -5 *T* [°C] $N_{\rm INP} = \lambda / V_{\rm droplet} = -\ln(1 - f_{\rm ice}) / V_{\rm droplet}$ λ -> for us: average number of ice active entities (INP) per droplet

*f*_{ice} -> fraction of frozen droplets from all droplets
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Normalization:

$$N_{\rm INP}$$
 = λ / $V_{\rm droplet}$ = - In (1 - $f_{\rm ice}$) / $V_{\rm droplet}$

Normalization can be done for:

- per volume of suspension (e.g., "K(T)" has been used)
 > V_{droplet} = volume of liquid in one droplet (e.g., for sea water samples)
- per volume of sampled air (airborne INP conc. N_{INP})
 -> V_{droplet} = volume of air collected into one droplet (all that I will show today)
- per surface area (surface site density n_s)
 -> V_{droplet} = surface area of material in one droplet



Normalization:

$$N_{\rm INP}$$
 = λ / $V_{\rm droplet}$ = - In (1 - $f_{\rm ice}$) / $V_{\rm droplet}$

Normalization can be done for:

per volume of suspension (e.g., "K(T)" has been used) 10^{-2} -> V_{droplet} = volume of liquid in one droplet (e.g., for sea water samples) per volume of sampled air (airborne INP conc. N_{INP}) [] ____10⁻³ $\rightarrow V_{\text{droplet}}$ = volume of air collected into one droplet (all that I will show today) per surface area (surface site density n_s) 10^{-4} -> V_{droplet} = surface area of material in one droplet, S_{droplet}) -20 -25 -15 -10 $n_s(T) =$ results from above equation after a Taylor but: *T* [°C] series expansion and is only valid for small f_{ice} Note that N_i is small compared with N_d (e.g., $N_i \approx 10^6 \text{ m}^{-3}$ and $N_d \approx 10^8 \text{ m}^{-3}$ at -28°C, corresponding to a frozen fraction f_i of about 1%), * as introduced in Niemand et al. (2012), A particlesurface-area-based parameterization of Immersion TROPOS 9 4.8.2020 @ online INP colloquium freezing on desert dust particles, JAS.



http://2.bp.blogspot.com/-IpuWzBRw3OY/VfQ1uo99E8I/AAAAAAABsg/jL8uHGMgZ2g/s1600/miller_world_physical_wall_mural_lg.jpg

<u>N_{INP} in Beijing during clean and polluted conditions</u>



concentrations unaffected by urban air pollution in Beijing, China, ACP.

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<u>N_{INP} in Beijing during clean and polluted conditions</u>



Chen et al. (2018), Ice nucleating particle concentrations unaffected by urban air pollution in Beijing, China, ACP.



<u>N_{INP} in Beijing during clean and polluted conditions</u>



- pollution in Beijing did not add INP that are ice active in the temperature range above -25°C

Chen et al. (2018), Ice nucleating particle concentrations unaffected by urban air pollution in Beijing, China, ACP.



annual series of filters from four Arctic stations

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filters collected airborne

samples from two ice cores

World Map

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 $http://2.bp.blogspot.com/-IpuWzBRw3OY/VfQ1uo99E8I/AAAAAAABsg/jL8uHGMgZ2g/s1600/miller_world_physical_wall_mural_lg.jpg$

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<u>*N*_{INP} from two Arctic glaciers</u>

two Arctic ice cores (Greenland and Svalbard) spanning the years 1478 to 1989

-> no increase in N_{INP} active down to -20°C

N_{INP} at -20°C 10¹ 107 a 10^{0} 106 10^{-1} 105 10-2 104 10-3 N_{INP} at -15°C 10¹ 107 b N_{INP} (L⁻¹ water) 10⁰ NINP 106 10^{-1} 0 105 10-2 air) 104 10-3 10³ 10^{-4} N_{INP} at -10°C 10⁰ 10⁶ 10^{-1} 105 10^{-2} 104 10-3 10³ 10^{-4} 10² 10-5 10¹ 10^{-6} 1500 1600 1700 1800 1900 2000 Age CE EUROCORE, multiyear samples Lomo09, multiyear samples Lomo09, subyear samples

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Hartmann et al. (2019), Variation of ice nucleating particles in the European Arctic over the last centuries, GRL.



Wex et al. (2019), Annual variability of ice nucleating particle concentrations at different Arctic locations, ACP.

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high concentrations in summer, origin terrestrial or marine

more indications, that human pollution does not add INP (active > -20°C)



Wex et al. (2019), Annual variability of ice nucleating particle concentrations at different Arctic locations, ACP.

high concentrations in summer, origin terrestrial or marine

more indications, that human pollution does not add INP (active > -20°C)

Borys (1983), The effects of long-range transport of air pollutants on Arctic cloudactive aerosol, PhD thesis.

Bigg & Leck (2001), Cloud-active particles over the central Arctic Ocean, JGR.

Creamean et al. (2018), Marine and terrestrial influences on ice nucleating particles during continuous springtime measurements in an Arctic oilfield location, ACP.

Tobo et al. (2019), Glacially sourced dust as a potentially significant source of ice nucleating particles, Nat. Geosci.

Wex et al. (2019), Annual variability of ice nucleating particle concentrations at different Arctic locations, ACP.



intermezzo: biogenic INP (proteins)

heating samples to test for proteinaceous INP





high concentrations in summer, origin terrestrial or marine

more indications, that human pollution does not add INP (active > -20°C)

biogenic

for VRS:

unheated and heated (95°C for 1 h) samples from INDA and the unheated samples from LINA



Wex et al. (2019), Annual variability of ice nucleating particle concentrations at different Arctic locations, ACP.



high concentrations in summer, origin terrestrial or marine

more indications, that human pollution does not add INP (active > -20°C)

biogenic

Zeppenfeld et al. (2019), Glucose as a potential chemical marker for ice nucleating activity in Arctic seawater and melt pond samples, ES&T.

Šantl-Temkiv et al. (2019), Biogenic sources of Ice Nucleation Particles at the high Arctic site Villum Research Station, ES&T.

Wex et al. (2019), Annual variability of ice nucleating particle concentrations at different Arctic locations, ACP.

for VRS:

unheated and heated (95°C for 1 h) samples from INDA and the unheated samples from LINA





airborne Arctic N_{INP}

at least in March over polynyas, INP (likely biogenic) originated locally from the ocean area

Low INP sample 10⁰ Low INP field blank 25.3.2018 (blank) 30,3,2018 (blank) 31.3.2018 (blank) 25,3,2018 30,3,2018 10-1 31.3.2018 INP concentration (L⁻¹) 10-2 10-3 10⁻⁴____30 -25 -20 -15-10 -5 Temperature in °C

Blank fit:

Boundaries from Petters & Wright 2015

 $N_{INP} = \exp(T * -0.51 + -14.60)$

Hartmann et al. (2020), Wintertime airborne measurements of ice nucleating particles in the high Arctic: a hint to a marine, biogenic source for Ice Nucleating Particles, GRL.

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intermezzo: Arcic Sea Ice Extent



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charctic-

ice-graph/

interactive-sea-

<u>N_{INP} in the Arctic</u>



- more indications, that human pollution does not add INP (active > -20°C)

- two ice cores

- high concentrations in summer

- origin terrestrial or marine

- at least in March over polynyas, INP originated locally from the ocean area

Hartmann et al. (2019), Variation of ice nucleating particles in the European Arctic over the last centuries, GRL.

Hartmann et al. (2020), Wintertime airborne measurements of ice nucleating particles in the high Arctic: a hint to a marine, biogenic source for Ice Nucleating Particles, GRL. Šantl-Temkiv et al. (2019), Biogenic sources of Ice Nucleation Particles at the high Arctic site Villum Research Station, ES&T.

Wex et al. (2019), Annual variability of ice nucleating particle concentrations at different Arctic locations, ACP.



filters from sea level (CVAO) and a ~ 700m high mountain (MV) on Sao Vicente, Cabo Verde

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GREENLAND Kalaallit Nuna



Welti et al. (2018), Concentration and
variability of ice nuclei in the subtropical
maritime boundary layer, ACP.

Gong et al. (2020), Characterization of aerosol particles at Cape Verde close to sea and cloud level heights - Part 2: ice nucleating particles in air, cloud and seawater, ACP.





Gong et al. (2020), Characterization of aerosol particles at Cape Verde close to sea and cloud level heights - Part 2: ice nucleating particles in air, cloud and seawater, ACP.





and cloud level heights - Part 2: ice nucleating particles in air, cloud and seawater, ACP.

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-20 Temperature [°C]

-15

-25

-10

CVAO 1596 CVAO 1596 heated CVAO 1641 CVAO 1641 heated CVAO 1643 CVAO 1643 heated Welti et al. (2018) 0.1 N_{INP} [# L_{air} ⁻¹, STP] 10'0 0.001 -25 -20 -15 -10 -5 Temperature [°C]

biogenic INP were present

Gong et al. (2020), Characterization of aerosol particles at Cape Verde close to sea and cloud level heights - Part 2: ice nucleating particles in air, cloud and seawater, ACP.



at Cape Verde, the marine boundary layer was well mixed

N_{INP} in air and cloud water fit well

compared by scaling with CCNconcentrations and an assumed droplet size



Gong et al. (2020), Characterization of aerosol particles at Cape Verde close to sea and cloud level heights - Part 2: ice nucleating particles in air, cloud and seawater, ACP.







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- most (> 70%) of all INP were supermicron in size
- biogenic INP were present
- marine boundary layer was well mixed
- $N_{\rm INP}$ in air and cloud water fit well
- INP from the ocean likely only contributed a small fraction of all INP

Gong et al. (2020), Characterization of aerosol particles at Cape Verde close to sea and cloud level heights - Part 2: ice nucleating particles in air, cloud and seawater, ACP.



the picture I have in my head

human pollution does not add
 INP (active > -25°C)

- supermicron INP are important in the atmosphere

high ice activity at high ice
 nucleation temperatures is
 (often) related to biogenic INP

- INP origin terrestrial or marine

-> it is often assumed that terrestrial sources are overwhelming, but the source strength of the ocean is still unknown, as well as biogenic contributions from terrestrial areas



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